

# Performance of linear mixed models and random forests for spatial prediction of Soil pH.

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31 **ABSTRACT**

32 Digital soil maps describe the spatial variation of soil and provide important information on  
33 spatial variation of soil properties which provides policy makers with a synoptic view of the  
34 state of the soil. This paper presents a study to tackle the task of how to map the spatial variation  
35 of soil pH across Zambia. This was part of a project to assess suitability for rice production  
36 across the country. Legacy data on the target variable were available along with additional  
37 exhaustive environmental covariates as potential predictor variables. We had the option of  
38 undertaking spatial prediction by geostatistical or machine learning methods. We set out to  
39 compare the approaches from the selection of predictor variables through to model validation,  
40 and to test the predictors on a set of validation observations. We also addressed the problem of  
41 how to robustly validate models from legacy data when these have, as is often the case, a  
42 strongly clustered spatial distribution. The validation statistics results showed that the empirical  
43 best linear unbiased predictor (EBLUP) with the only fixed effect a constant mean (ordinary  
44 kriging) performed better than the other methods. Random forests had the largest model-based  
45 estimates of the expected squared errors. We also noticed that the random forest algorithm was  
46 prone to select as “important” spatially correlated random variables which we had simulated.

47 **Keyword:** Linear Mixed Models; REML-EBLUP; Random forests, Spatial prediction of soil  
48 pH.

49 **1. INTRODUCTION**

50 Soil maps describe the spatial variation of soil types and provide important information on spatial  
51 variation of soil properties (Kempen et al., 2010). Mapping of soil properties is important as it  
52 provides policy makers with a synoptic view of the state of the soil, and agricultural stakeholders  
53 with information about where soil problems might occur (Lark et al., 2019). Soil maps are

54 generated using various soil mapping methods which can be divided into conventional and  
55 pedometric approaches (Kienast-Brown et al., 2010; Hengl, 2003).

56 Conventional soil survey represents soil variation in terms of profile classes and corresponding  
57 map legend units. It can provide a basis for spatial prediction of soil properties and may also serve  
58 as a structure for recording substantial information on soil management and for systematizing  
59 knowledge of the distribution of soils in the landscape. Conventional approaches were based  
60 largely on manual processes which are costly and time consuming (Kienast-Brown et al., 2010)  
61 mainly because of long fieldwork periods (Moonjun et al., 2010). Pedometric approaches are based  
62 on the application of mathematical and statistical methods for the primary purpose of predicting  
63 the values of soil properties where these have not been observed directly (McBratney et al., 2000).  
64 A well-established statistical approach to doing this is the application of model-based geostatistics  
65 (Stein, 1999; Diggle and Ribeiro, 2007). In this approach the variation of the soil is represented in  
66 a linear mixed model (LMM) as a combination of fixed effects (which may be a constant unknown  
67 mean, or a function of predictive covariates such as remote sensor data), and random effects,  
68 including Gaussian random fields which exhibit spatial correlation. The parameters of the LMM  
69 model can be estimated by Residual Maximum Likelihood (REML) method developed by  
70 Patterson and Thompson (1971), which allows parameters of the random effects to be estimated  
71 with small bias arising from uncertainty in the fixed effects (Kitanidis, 1987; Swallow and  
72 Monahan 1984; Zimmerman and Zimmerman, 1991; Lark and Cullis, 2004) . When the model is  
73 fixed, values of the soil property at unsampled sites can be obtained by the empirical best linear  
74 unbiased predictor (EBLUP) (Stein, 1999; Lark et al., 2006; Lark and Webster, 2006; Minsay and  
75 McBratney, 2007).

76 There has been a growing interest in the potential of machine learning methods (e.g. Breiman,  
77 2001) as an alternative to statistical modelling for spatial prediction of soil properties (Hengl et al.,  
78 2015; Behrens and Scholten, 2007). The main difference between geostatistical approaches and

79 random forest is that geostatistics is based on a statistical model. This provides a basis for formal  
80 inference about the validity of the model (including the task of selecting which covariates to use in  
81 prediction), and for producing a prediction distribution at unsampled sites of interest. One may  
82 then derive point predictions from this distribution (typically the mean), and measures of  
83 uncertainty. On the other hand, machine learning methods such as random forests, are predictive  
84 tools applied to identify empirical relationships between the target variable in a training data set  
85 and associated predictive covariates and to extrapolate these to unsampled sites. With no model  
86 there can be no formal inference, but empirical approaches, based on internal cross-validation are  
87 used, for example, to evaluate the evidence that a particular variable is predictive. One particular  
88 strength of the geostatistical approach is that the estimation of coefficients for predictor variables,  
89 and inferences about them, are based on a model of the spatial dependence of the random variation.  
90 This accounts for the fact that data which are strongly spatially clustered are likely to be correlated,  
91 and so do not provide independent evidence to support the fitted model.

92 In the study reported here our objective was to assess approaches for digital mapping of soil pH at  
93 national scale across Zambia to support evaluation of land potential for rice production. Legacy  
94 data on soil pH were available from a previous national survey. As with many such surveys, this  
95 followed a two-stage design, and so the observations were spatially clustered. In addition we had  
96 access to various exhaustive environmental covariates which could be regarded as potential  
97 predictor variables for soil pH.. We compared different forms of linear mixed model, and  
98 prediction with the random forest using a validation subset of the data. Prediction errors were  
99 evaluated at the validation locations by comparing predictions with observed values. The selection  
100 of the validation subset, and the quantification of the uncertainty from the observed prediction  
101 errors had to take account of the spatial clustering of the observations in the legacy data. Because  
102 of this clustering, no subset could be regarded as independent random observations.

## 103 **2. THEORY**

104 **2.1 Linear Mixed Model**

105 The theory of residual maximum likelihood (REML) in combination with the empirical best  
 106 linear unbiased predictor (EBLUP) for spatial interpolation has been illustrated and described  
 107 in detail by Lark et al., 2006. The LMM takes the form

$$108 \quad \mathbf{z} = \mathbf{M}\boldsymbol{\beta} + \mathbf{S}\boldsymbol{\eta} + \boldsymbol{\varepsilon}, \quad (1)$$

109

110 where  $\mathbf{z}$  is a set of observations of the random variable at sampled locations,  $\mathbf{M}$  is the design  
 111 matrix of fixed effects, which could include covariates such as topographic attributes,  $\boldsymbol{\beta}$  is the  
 112 vector of the fixed effects parameters or regression coefficients,  $\mathbf{S}$  is the design matrix of  
 113 random effects (which is an identity matrix unless analytical duplicate observations are  
 114 included),  $\boldsymbol{\eta}$  is a random effect, a Gaussian random variable which has a mean of zero and, in  
 115 the spatial setting, a covariance matrix which expresses spatial dependence,  $\boldsymbol{\varepsilon}$  is an  
 116 independently and identically distributed Gaussian residual of mean zero and variance  $\sigma^2$ .  
 117 These two random components have a joint distribution

$$118 \quad \begin{bmatrix} \boldsymbol{\eta} \\ \boldsymbol{\varepsilon} \end{bmatrix} \sim \mathcal{N} \left( \begin{bmatrix} \mathbf{0} \\ \mathbf{0} \end{bmatrix}, \begin{bmatrix} \sigma^2 \boldsymbol{\xi} \mathbf{G} & \mathbf{0} \\ \mathbf{0} & \sigma^2 \mathbf{I} \end{bmatrix} \right), \quad (2)$$

119

120 where  $\mathbf{I}$  is the identity matrix and  $\mathbf{G}$  is the correlation matrix of the random effect  $\boldsymbol{\eta}$ . Element  
 121  $[i,j]$  of  $\mathbf{G}$ , at locations  $\mathbf{x}_i$  and  $\mathbf{x}_j$  depends only on the interval in space between them under an  
 122 assumption of second-order stationarity. The lag vector  $\mathbf{x}_i - \mathbf{x}_j$ , under the assumption of  
 123 isotropy, depends only on the scalar part of this vector, the lag distance and so

$$124 \quad \mathbf{G}[i,j] = \rho(|\mathbf{x}_i - \mathbf{x}_j|; \alpha), \quad (3)$$

125

126 where  $\rho(h; \alpha)$  is a correlation function of lag distance  $h$  with spatial parameters  $\alpha$  which control  
 127 how the correlation decreases with increasing distance. The term  $\xi$  is the ratio of the variance  
 128 of the random effect  $\boldsymbol{\eta}$  to  $\sigma^2$ , the variance of the residual term.

129 The residuals depend on the fixed effects parameters  $\boldsymbol{\beta}$  in the model, and in ordinary maximum  
 130 likelihood estimation the uncertainty in the estimates of the fixed effects parameters biases the  
 131 estimates of the random effects parameters. To avoid this, we use residual maximum likelihood  
 132 (REML) which is based on the principle that a new random variable, independent of the fixed  
 133 effects, is computed by projecting the original data  $\mathbf{z}$  into a residual space where the fixed  
 134 effects can be filtered out (Chai et al., 2008). The log likelihood of the new random variable  
 135 which we now call the residual log-likelihood because its independent of fixed effects can be  
 136 expressed as;

$$137 \quad \ell_R(\sigma^2, \xi, \alpha | \mathbf{z}) = -\frac{1}{2} \{ \log |\mathbf{H}| + \log |\mathbf{M}^T \mathbf{H} \mathbf{M}| + (n-p)\sigma^2 + \frac{1}{\sigma^2} \mathbf{z}^T (\mathbf{I} - \mathbf{W} \mathbf{C}^{-1} \mathbf{W}^T) \mathbf{z} \}, \quad (4)$$

$$138 \quad \text{Where } \mathbf{H} = \xi \mathbf{M} \mathbf{G} \mathbf{Z}^T + \mathbf{I}, \mathbf{W} = [\mathbf{M}, \mathbf{S}] \text{ and } \mathbf{C} = \begin{bmatrix} \mathbf{M}^T \mathbf{M} & \mathbf{M}^T \mathbf{S} \\ \mathbf{S}^T \mathbf{M} & \mathbf{S}^T \mathbf{S} + \xi^{-1} \mathbf{G}^{-1} \end{bmatrix}.$$

139 Once the covariance parameters  $\sigma^2, \xi, \alpha$  have been estimated by REML, they are used to  
 140 compute the estimated covariance matrix at sampled points. With the estimated covariance  
 141 matrix computed, the estimated fixed effects parameter,  $\hat{\boldsymbol{\beta}}$ , and predicted random effects,  $\hat{\boldsymbol{\eta}}$ , are  
 142 then computed by solution of mixed model equation:

$$143 \quad \mathbf{C} \begin{bmatrix} \hat{\boldsymbol{\beta}} \\ \hat{\boldsymbol{\eta}} \end{bmatrix} = \begin{bmatrix} \mathbf{M}^T \mathbf{z} \\ \mathbf{S}^T \mathbf{z} \end{bmatrix} \quad (5)$$

144 With the covariance matrix for the error of the estimates being:

$$145 \quad \text{Cov} \begin{bmatrix} \hat{\boldsymbol{\beta}} - \boldsymbol{\beta} \\ \hat{\boldsymbol{\eta}} - \boldsymbol{\eta} \end{bmatrix} = \sigma^2 \mathbf{C}^{-1}. \quad (6)$$

146 Now that the covariance matrix and the fixed effects parameters have been estimated, they are  
 147 used in EBLUP to predict the soil property ,  $\tilde{z}_p$  , at unsampled locations:

$$148 \quad \tilde{z}_p = \mathbf{M}_p^T \hat{\boldsymbol{\beta}} + \tilde{\boldsymbol{\eta}}_p = \mathbf{M}_p^T \hat{\boldsymbol{\beta}} + \mathbf{g}_{o,p}^T \mathbf{G}^{-1} \tilde{\boldsymbol{\eta}}, \quad (7)$$

149 where  $\mathbf{M}_p$  is the design matrix for the prediction sites,  $\mathbf{g}_{o,p}$  is a vector computed from the  
 150 covariance matrix of  $\boldsymbol{\eta}$  with the  $\boldsymbol{\eta}_p$  values at the unsampled locations ( $\text{Cov}[\boldsymbol{\eta}, \boldsymbol{\eta}_p] = \xi \sigma^2 \mathbf{g}_{o,p}$ ).

151 The variance of the prediction errors,  $\text{Var}[\tilde{z}_p - z_p]$ , which accounts for the uncertainty in  
 152 predicting the fixed effects and uncertainty in predicting the random effects is expressed as:

$$153 \quad \text{Var}[\tilde{z}_p - z_p] = \sigma^2 \left\{ [\mathbf{M}_p, \mathbf{g}_{o,p}^T \mathbf{G}^{-1}]^T \mathbf{C}^{-1} [\mathbf{M}_p, \mathbf{g}_{o,p}^T \mathbf{G}^{-1}] + \xi (\mathbf{g}_{p,p} - \mathbf{g}_{o,p}^T \mathbf{G}^{-1} \mathbf{g}_{o,p}) + 1 \right\}. \quad (8)$$

154 There are many variables that researchers can use as fixed effects in linear mixed models for  
 155 spatial prediction of soil properties. However, it is unwise to include variables without regard  
 156 for evidence that they are of predictive value, the inclusion of predictors unrelated to the target  
 157 variable may inflate the prediction error variance. To avoid this, variable selection is an  
 158 important step in model development. One approach to the problem is to base the inclusion or  
 159 rejection of a predictor based on a hypothesis test in the LMM framework (e.g. by a log-  
 160 likelihood ratio test) (Verbeke and Molenberghs, 2000). To reduce the risk of including excess  
 161 predictors because of multiple hypothesis testing, one may use false discovery rate control  
 162 (Lark, 2017). The false discovery rate (FDR) is the probability that a null hypothesis is true,  
 163 given that it has been rejected. False discovery rate control can reduce the power to detect real  
 164 predictors, and Lark, (2017) demonstrated how this problem can be reduced, while maintaining  
 165 FDR control, by the method of alpha investment (Foster and Stine, 2008). This entails an initial  
 166 ordering of the predictors starting with the one which, a priori (and not based on inspection of  
 167 the data) is thought most likely to be of predictive value and adding in predictors in declining  
 168 order of expected predictive power. In this approach the power to detect a predictor is increased

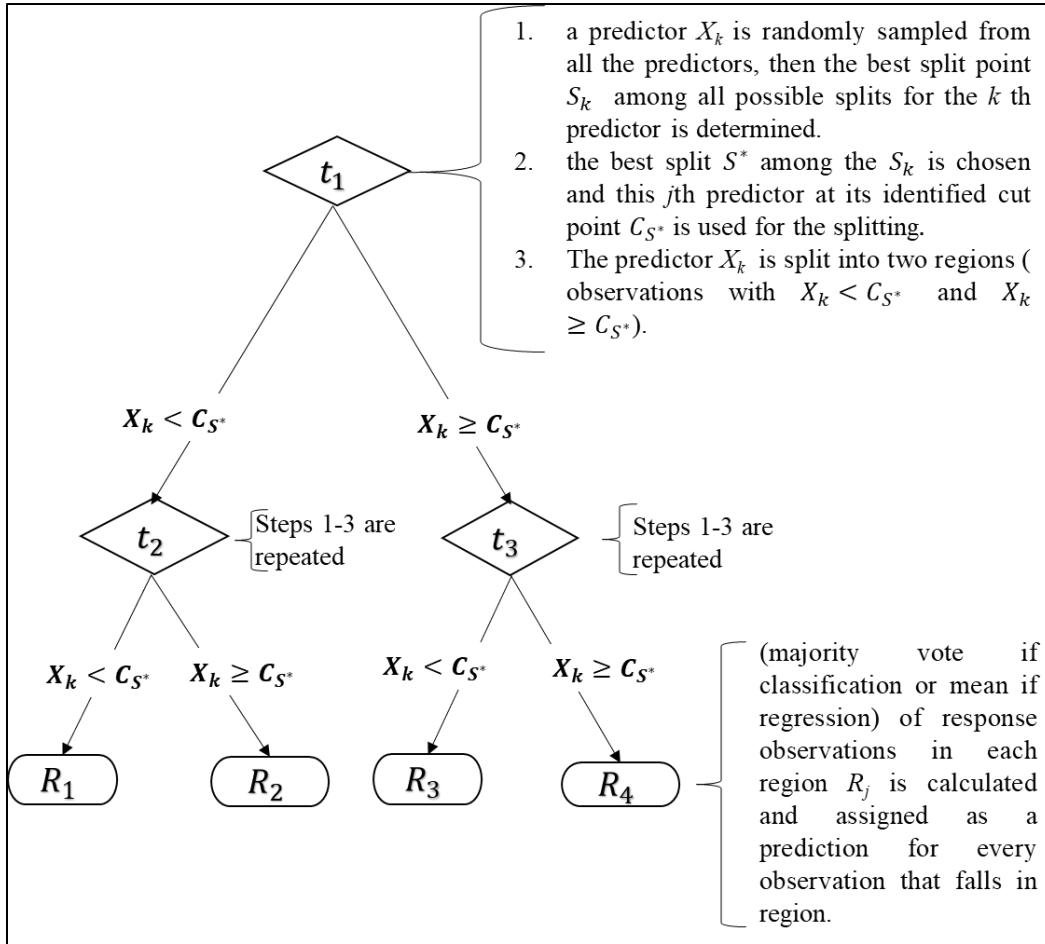
169 by the rejection of null hypotheses early in the sequence while maintaining control of FDR.  
170 This approach has been used elsewhere for spatial prediction (Gashu et al., 2020).

171 A disadvantage of the LMM approach is that it assumes that the fixed effects are linear in the  
172 parameters. Such a model can represent complex and non-linear relations between soil  
173 properties and predictors, for example through the use of polynomial terms in the predictor  
174 variables, or spline basis functions, but there has been increasing interest in more flexible  
175 methods to predict soil properties from covariates, in particular the machine learning method  
176 known as the random forest.

## 177 **2.2 Random Forests**

178 The random forest is an ensemble tree-based method that combines multiple decision trees  
179 (classification or regression) to give a prediction (Breiman, 2001). A decision tree is an  
180 algorithm that involves recursive partitioning of data into several simple regions using a series  
181 of splitting rules. It is called a decision tree because these series of splitting rules can be  
182 summarised into an upside-down tree structure as illustrated in Figure 1. Figure 1 shows a  
183 structure made up of predictors ( $X_1, X_2, \dots, X_k$ ) which are split into  $J$  distinct and non-  
184 overlapping regions ( $R_1, R_2, \dots, R_j$ ) at test node  $t$ , and the mean of the response values for the  
185 training observations in each region  $R_j$  is calculated and assigned as a prediction for every  
186 observation that falls in region  $R_j$  (James et al., 2013). When growing a decision tree, the  
187 following steps are taken; (1) at each test node  $t$ , a predictor  $X_k$  is randomly sampled from all  
188 the predictors, then the best split point  $S_k$  among all possible splits for the  $k$ th predictor is  
189 determined; (2) the best split  $S^*$  among the  $S_k$  is chosen and this  $j$ th predictor at its identified  
190 cut point  $C_{S^*}$  is used for the splitting at test node  $t$ . (3) The predictor  $X_k$  is split into two regions  
191 ( $X_k < C_{S^*}$  and  $X_k \geq C_{S^*}$ ) at test node  $t$ . Steps 1-3 are repeated on all  
192 descendant nodes to grow a tree  $\hat{f}(x)$  (Archer and Kimes, 2008; James et al., 2013).





193

194 *Figure 1: illustration of a decision tree*

195 One major limitation with decision trees is that using only one tree for prediction, results in  
 196 highly unstable predictions, a small change in the data can result into a large change in the final  
 197 estimated tree. To improve the performance of decision trees, Breiman (1996) introduced an  
 198 algorithm called Bagging, also known as bootstrap aggregation which takes repeated  
 199 (bootstrap) samples (where  $B$  is the number of bootstrap samples) from training set with  
 200 replacement and builds a total of  $B$  trees ( $\hat{f}^1(x), \hat{f}^2(x), \dots, \hat{f}^B(x)$ ) which the average of all the  
 201 prediction trees  $\hat{f}_{bag}(x)$  is calculated:

202

$$\hat{f}_{bag}(x) = \frac{1}{B} \sum_{b=1}^B \hat{f}^b(x) \quad (9)$$

203 One disadvantage of bagging is that a single predictor may dominate all trees in the bag,  
204 meaning that their outputs are strongly correlated. As a result of this the reduction in variance  
205 from the use of multiple trees is very limited (James et al., 2013). To address this, Breiman  
206 (2001) developed a random forest algorithm which is an improvement of bagging. Like  
207 bagging, random forest also takes repeated (bootstrap) samples from the training data and  
208 builds  $B$  decision trees. But in the case of random forest, when building these trees, to avoid  
209 using one strong predictor for all bagged trees, at every test node  $t$ , when splitting, every bagged  
210 tree is forced to consider only a random subset of predictors by randomly sampling a fresh  $m$   
211 predictors from a set of  $k$  predictors, and the split is only allowed to use one of these  $m$   
212 predictors. For regression trees, the number of  $m$  predictors considered at each split is  
213 approximately the total number of predictors divided by three ( $m \approx k/3$ ) and for classification  
214 trees,  $m \approx \sqrt{k}$ . Because of this, random forest results in many uncorrelated trees which give a  
215 large reduction in variance when averaged.

216 The random forest algorithm has three important outputs. These are the out-of-bag Mean  
217 Squared Error (OOB Mean Squared Error), the out-of-bag R-squared and the variable  
218 importance. The RF model does not use all the data for building the tree. In each bootstrap  
219 training set, about one-third of the data are left out. The data that are left out when building the  
220 trees is called out-of-bag (OOB) data and after the trees are grown, the OOB data are used as  
221 test set to measure the strength (OOB Mean Squared Error) and correlation (OOB R-squared)  
222 of the model. In short, random forest has an inbuilt cross-validation. Variable importance is  
223 defined as the increase in prediction error when OOB data for that variable is randomly  
224 permuted while all others are left unchanged (Liaw, 2002). It analyses the contributions of each  
225 predictor to the overall results (Breiman, 2001). The algorithm randomly permutes the predictor  
226  $X_m$  several times, breaking its original association with the response variable and assesses the  
227 relevance of the predictor by using the permuted predictor together with the other unpermuted

228 predictors to predict the response variable for the out-of-bag observations giving the difference  
229 in prediction accuracy before and after permuting. The result is a vector of importance measures  
230 for each predictor equivalent to the number of permutations. The algorithm then computes a p-  
231 value as a measure of the evidence that a variable is predictive (Strobl et al., 2007; Altmann et  
232 al., 2010). This permutation p-value is the probability of observing a permuted model (from  
233 the several number of permutations) that is equal to or better than the unpermuted model  
234 (Cummings et al., 2004).

235 Equation (7) presents the E-BLUP from a linear mixed model. The second term on the right-  
236 hand side corresponds to the spatial interpolation of the correlated random effect in the model.  
237 In this way the E-BLUP combines a regression-type prediction based on the predictor variables  
238 with a spatial interpolation component. As described above the random forest predicts a soil  
239 property from the predictor variables only, making no use of spatial dependence through  
240 interpolation. An attempts has been made to include spatially weighted local observations in  
241 prediction by random forests by including coordinates as predictors and using weighted buffer  
242 distances (Hengl et al., 2018), neighbouring observations and their distances to the prediction  
243 location were used by Sekulić et al., (2020). Li et al., (2011) and Viscarra Rossel et al., (2014)  
244 combined random forest with kriging, just like in regression-kriging, by calculating the random  
245 forest residuals and then kriged them to all prediction positions and then added to the results of  
246 the prediction positions. .

247 As described above, inferences in the random forest approach are based on an internal cross-  
248 validation procedure. This might lead to overoptimistic conclusions about a random forest  
249 model, or about the value of a particular predictor if observations from the same clusters appear  
250 in the OOB sample and in the data used to develop the trees. That is because the observations  
251 within a cluster can be expected to be strongly correlated, and so the validation of a model fitted

252 to data on strongly correlated observations will give an unduly optimistic impression of the  
253 model's capacity to predict at an independent location.

### 254 **3. CASE STUDY**

#### 255 **3.1. Data**

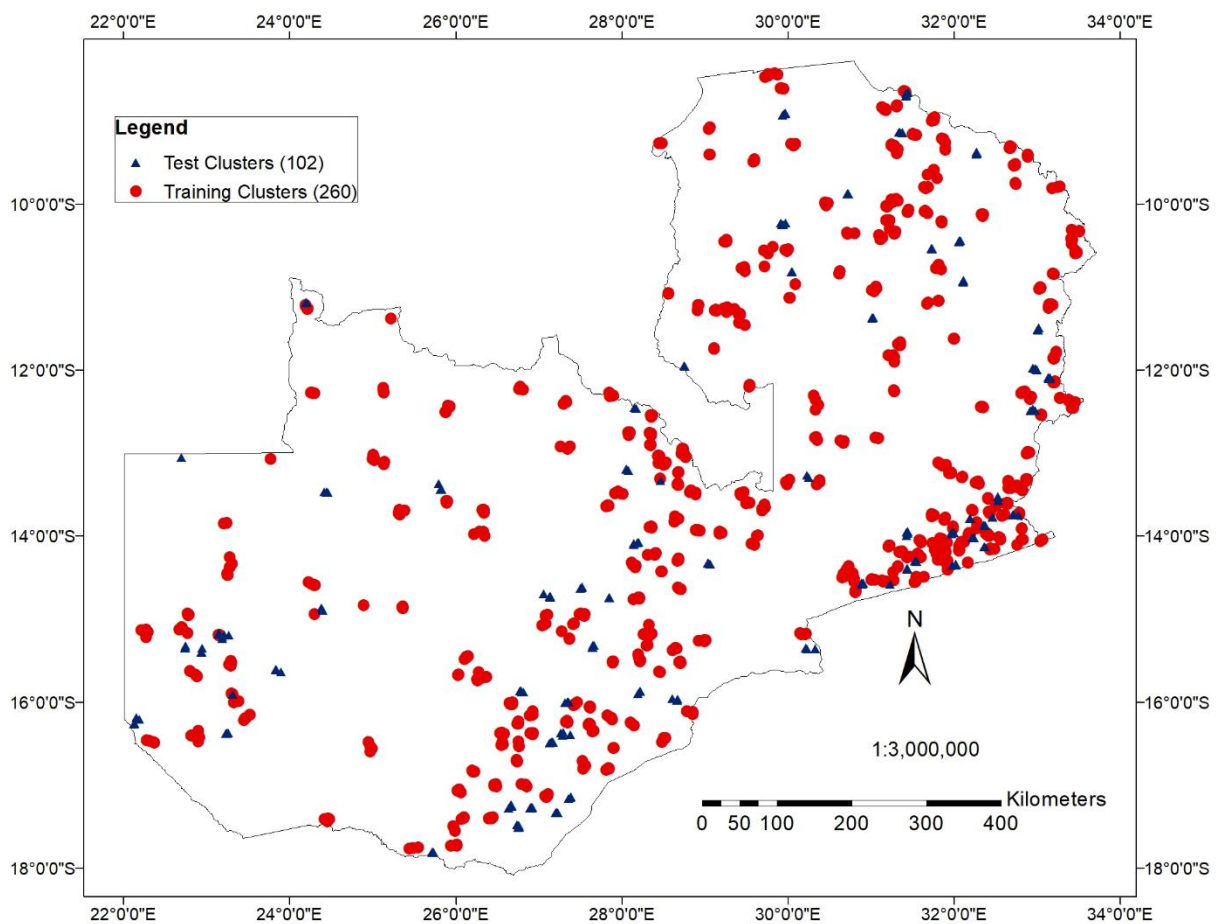
##### 256 **Soil data**

257 This case study uses Rural Agricultural Livelihoods Survey (RALS) of 1713 soil pH data  
258 collected by Indaba Agricultural Policy Research Institute (IAPRI) in collaboration with  
259 Central Statistical Office (CSO) and Ministry of Agriculture. The sampling frame for the RALS  
260 2012 survey was based on the 2010 Census of Housing and Population (CSO/MAL/IAPRI,  
261 2015). Full detail of the stratified two-stage sampling design is provided by (CSO 2012). Four  
262 households were randomly selected in each Standard Enumeration Area (SEA) and soil samples  
263 were collected from the largest maize field. A composite of 10–20 sub-samples of soil collected  
264 throughout each field and each sub-sample was a composite of equal parts soil in the 0–10cm  
265 and 10–20cm depth horizons. Full details on the soil collection and laboratory analysis for soil  
266 pH (determined for a soil suspension in  $\text{CaCl}_2$  with a standard pH meter) are provided by Burk  
267 et al., (2019) and Chapoto et al., (2016). The spatial prediction of soil pH for Zambia using this  
268 data has been studied by Chapoto et al., (2016) who only used ordinary kriging for the  
269 prediction.

270 Data cleaning involved removal of spurious values in the x and y coordinates. The need for  
271 this was indicated when the raw data were first plotted, showing points lying outside the borders  
272 of Zambia. The mean coordinates of all households were computed in each Standard  
273 Enumeration Area (SEA) centroid, and then the households were removed from the data set if  
274 the notional distance to the SEA centroid exceeded 10km. This threshold value was decided  
275 after discussion with IAPRI staff about plausible values for the distance between a village in  
276 the EA and the centroid. After data cleaning, a total of 1202 soil samples were used for analysis.

277 The sampling pattern for the RALS survey was not designed for spatial interpolation of soil.  
278 Due to the sampling pattern, the data is strongly clustered at the level of SEAs (the SEAs are  
279 the clusters) with a total of 362 clusters. For this reason, splitting of the dataset into training set  
280 (80%) and test set (20%) was done at cluster level (the 362 clusters were split into 260 (80%)  
281 for training and 102 (20%) for validation). Figure 2 Shows the training and test clusters with  
282 the red solid dots being the training clusters and the blue solid triangles being the test clusters.

283



285 *Figure 2: Cluster locations for the RALS 2012 soil data. red solid dots being the training*  
286 *clusters used for spatial prediction of soil pH and the blue solid triangles being the test clusters*  
287 *left out for validation.*

## 288 Environmental Covariates

289 The environmental predictors available for use in this study were Soilclass, Landcover, mean  
 290 annual rainfall, elevation, slope, aspect, valley depth, LS-Factor (a combination of slope and  
 291 slope length, relative slope position (RSP), channel network base Level (CNBL) and  
 292 Normalized difference vegetation index (NDVI).

293 Soilclass information was obtained from the 1:1,000,000 scale exploratory soil map of Zambia  
 294 compiled by the *Zambian Ministry of Agriculture, Zambia Agricultural Research Institute*  
 295 (*ZARI*) - Soil Survey Section in 1991(*GRZ*, 1991) and then digitized to raster format. Map  
 296 units are allocated to suborders of the *FAO-UNSECO* classification as used in the Third Draft  
 297 of the legend to the *Soil Map of the World* (*Jahn et al*, 2006)). A total of 96 soil classes were  
 298 represented in the data available for model development, but these do not comprise all the  
 299 classes on the map of Zambia, and so some generalization is required to develop models for  
 300 spatial prediction. We, therefore, reduced the number of classes, by aggregating the classes  
 301 from suborder to order level, this reduced the number of classes to 18 and all the classes in the  
 302 prediction grid where represented in the training set. Land cover data for the years between  
 303 2000 and 2015 with spatial resolution of 300m were downloaded from the *European Space*  
 304 *Agency (ESA)*, (2017). The data presented a similar situation as that of soilclass with landcover  
 305 classes in the prediction sites not being represented in the training set. We also reduced the  
 306 number of landcover classes, by aggregating them as shown in Table 1.

307 *Table 1: Aggregated landcover classes based on ESA, (2017)*

<b>New Class</b>	<b>ESA Class</b>	<b>Description</b>
1	10	rainfed cropland
	20	irrigated or post-flooding cropland
	30	Mosaic cropland (>50%) / natural vegetation (tree
2	11	Herbaceous cover
	40	herbaceous cover) (>50%) / cropland (<50%)
	110	Mosaic herbaceous cover (>50%) / tree and shrub (<50%)
3	12	Tree or shrub cover

	100	Mosaic tree and shrub (>50%) / herbaceous cover (<50%)
4	50	closed to open (>15%), evergreen, broadleaved, tree cover
	60	closed to open (>15%), deciduous, broadleaved, tree cover
	61	closed (>40%), deciduous, broadleaved, tree cover
	62	open (15-40%), deciduous, broadleaved, tree cover
5	120	Shrubland
	122	Shrubland deciduous
6	130	Grassland
7	160	fresh or brackish water, flooded, tree cover
	170	saline water, flooded, tree cover
	180	fresh/saline/brackish water, flooded Shrub or herbaceous cover
8	190	Urban areas
9	200	Consolidated bare areas
	202	Unconsolidated bare areas

308

309 Mean annual rainfall data (averages from 1970 to 2000) with a spatial resolution of 1km was  
310 downloaded from WorldClim website (Fick and Hijmans, 2017). A 90-m resolution NASA  
311 Shuttle Radar Topography Mission (SRTM3) Digital elevation model (DEM) was downloaded  
312 from USGS (2019) and projected to WGS 84 UTM 35 S. The DEM was pre-processed by filling  
313 sinks using the fill sinks (Planchon/Darboux, 2001) tool in Saga GIS, and then elevation, slope,  
314 aspect, valley depth, LS-Factor (a combination of slope and slope length), relative slope  
315 position (RSP) and channel network base Level (CNBL) data was extracted from the DEM  
316 using basic terrain analysis tool in Saga GIS. MODIS land surface reflectance (MOD009GA  
317 V6) was downloaded from USGS (2019). After downloading the respective data sets, Quantum  
318 GIS was used to project the data sets to WS 84 UTM 35s and then converted to the Integrated  
319 Land and Water Information System (ILWIS) format. Then ILWIS was used to harmonise all  
320 the raster files to the same extent and cell size of 1km. Normalized difference vegetation index  
321 (NDVI) was extracted from the remote sensing images using the imageIndices of the  
322 soilassessment package for the R platform (Omuto, 2020).

### 323 **3.2. Spatial Prediction of soil pH**

324 Soil pH is an important chemical property of the soil that affects its fertility status. This is  
325 because the availability of most essential plant nutrients is influenced by the levels of pH in the  
326 soil (Jones, 2012). There are two principal processes that affect the levels of soil pH in the soil  
327 (1) the production of H<sup>+</sup> ions and (2) the loss of basic cations from the soil. Eleven variables  
328 were available to be considered as possible predictors for soil pH.

329 In section 2.1 we explained how variable selection for the LMM included false discovery rate  
330 control, to avoid over-fitting, with alpha-investment to improve the probability of retaining  
331 covariates which are predictive as predictor variables. The alpha-investment approach is most  
332 effective if the predictors can be ordered, a priori, with the one thought most likely to be  
333 predictive ranked first and so on as shown in Table 3. It must be emphasized that this ranking  
334 is based on prior considerations about the underlying process, and not on exploration of the  
335 data. We did this ordering of exhaustive environmental covariates based on how they influence  
336 the production of H<sup>+</sup> ions and the loss of basic cations from the soil. Rainfall was proposed as  
337 the most influential factor at national scale. Soils in environments with large annual rainfall  
338 tend to have relatively low pH due to reduced based saturation resulting from loss of basic  
339 cations by leaching (McCauley, et al., 2009; Brady and Weil, 2014). For this reason more acid  
340 soils are expected in the northern parts of Zambia (Agroecological Region Three) and in the  
341 south (Agroecological Region One) where annual rainfall is much smaller (Veldkamp et al.,  
342 1984; GRZ and UNDP, 2009). Soil class was ranked second because the soil classes represent  
343 variations in soil parent material, weathering and rejuvenation of land surfaces and development  
344 of the soils. The old, highly weathered plateau soils in the northern part of the country have lost  
345 most of the basic cations. The sandy soils in the western part are easily leached with little  
346 accumulation of basic cations. On the on the hand, the Karoo group materials in the valleys are  
347 rich in basic cations resulting in high pH values. After soil class we included topographic



348 variables Slope, Elevation and Valley Depth. These should reflect processes such as the  
349 movement of water which carries with it dissolved basic cations from steep slope to flat areas,  
350 and the rejuvenation of weathered land surfaces which entails the removal of old highly-  
351 weathered material to reveal material with a larger content of weatherable minerals. We then  
352 included Landcover and the Normalized Difference Vegetation Index (NDVI). These will  
353 reflect effects of land management practices, including agricultural inputs, and decisions on  
354 land use which may depend on how local pH limits crop performance. The NDVI will also  
355 reflect local vigor of vegetation growth, which may be pH limited. Finally we included some  
356 further topographic variables which may reflect differences in the soil-forming environment  
357 (length-slope factor, channel network base level, relative slope position and aspect.

358 The data points were first projected from WGS 1984 to WSG UTM 35s. A total of 19  
359 observations had duplicate coordinates, which were jittered by adding 100m to each of the  
360 coordinates for one site. An exploratory model was fitted to the data with all predictors  
361 included, using the `likfit` function of `geoR` package (Ribeiro and Diggle, 2001) with residual  
362 maximum likelihood (REML) as the likelihood method. The only output from this model which  
363 was examined were the residuals, for which summary statistics were calculated, and exploratory  
364 plots to check the plausibility of the assumption of normally distributed errors. In addition, the  
365 correlation model type (exponential or spherical) was identified for which the residual  
366 likelihood was largest, and this model was then used in all further analyses. During the  
367 sequence testing of hypothesis, first the null model,  $m_0$ , (with the only fixed effect a constant  
368 mean) was fitted with the `likfit` function and ML as the likelihood method. Then the next model,  
369  $m_1$ , with the first predictor in the sequence was fitted in the same way. The likelihood ratio was  
370 then calculated:

$$371 \quad L = 2(L_{m_1} - L_{m_0}), \quad (10)$$

372 where  $L_{m_1}$  is the likelihood for model  $m_1$  and  $L_{m_0}$  is the likelihood for the null model. If the  
373 null model is correct, then the asymptotic distribution of  $L$  is  $\chi^2$  with degrees of freedom equal  
374 to the number of additional parameters in model  $m_1$  by comparison with  $m_0$ . If  $L$  provided  
375 evidence to reject the null model with  $P < 0.05$ , then the additional predictor in model  $m_1$  was  
376 retained. The second predictor in the list was then considered. When all predictors had been  
377 examined the  $P$ -values at each step were reassessed in sequence using alpha-investment as  
378 described by Lark (2017) and controlling the FDR at 0.05. Details of this approach are provided  
379 by Lark (2017), but in summary successive tests are made against a threshold  $P$ -value which  
380 depends on a quantity, the alpha-wealth, which is either augmented when null hypotheses are  
381 rejected or augmented when they are rejected. If the hypotheses are ordered so that the variables  
382 which, a priori, are expected to be good predictor variables are considered early, then this alpha  
383 investment method increases the probability of selecting predictive covariates while controlling  
384 FDR.

385 After variable selection, the `likfit` function of `geOR` package in R with REML as the likelihood  
386 method was used to fit two linear mixed models. One with the selected predictors as fixed  
387 effects (Kriging with an external drift) and the other with a constant mean as fixed effect  
388 (ordinary kriging). The E-BLUP prediction for both models was then calculated at the  
389 validation points.

390 The `ranger` function of `ranger` package (Wright and Ziegler, 2017) was used to fit the random  
391 forest model. Because random forest has an inbuilt variable selection that occurs within the  
392 model by randomly selecting variables to be used at splitting nodes, two models were fitted,  
393 one with all variables and the other with the two variables that were selected during the alpha-  
394 investment variable selection procedure. In this study, we use the `ranger` package in R to  
395 compute the permutation variable importance according to Altmann et al., (2010).

396 When predicting soil properties in space, the random forest algorithm can find apparently  
397 predictive relationships between the target variable and arbitrary spatial variables (such as  
398 digital images of human faces) when these are presented as candidate predictor variables  
399 alongside covariates which pedometricians might reasonably expect to be predictive of soil  
400 properties, (Wadoux et al., 2020). This shows that pattern recognition should not be equated to  
401 knowledge discovery. It may also suggest that the random forest algorithm is prone to  
402 overfitting, as a result of which its predictions at independent locations may be unreliable. To  
403 investigate this effect, we generated entirely random spatially autocorrelated candidate  
404 predictor variables, independent of our data, which we call null predictors. We used six spatially  
405 correlated but mutually independent null predictors, specifying a spherical variogram with a  
406 distance parameter of 100 km, nugget variance of 0 and correlated variance of 1 for each. We  
407 used the function `RFsimulate` from the `RandomFields` package for R (Schlather et al., 2015) to  
408 simulate values of these null predictors at the calibration locations. We then used the `ranger`  
409 package (Wright and Ziegler, 2017) to fit a random forest model with all predictors including  
410 the null variables as predictors and then computed the permutation variable importance  
411 according predictor p-values from the model result.

412 To examine the possibility of improving RF predictions by an additional kriging step (following  
413 Li et al., 2011 and Viscarra Rossel et al., 2014), residuals of the models at training points were  
414 derived (subtracting the predicted values from the observed values) and then a variogram was  
415 fitted to the residuals using `likfit` with a constant mean as the only fixed effect. The evidence  
416 for spatial dependence in the residuals was assessed by comparing the Akaike Information  
417 Criterion (AIC) for the fitted model and for a non-spatial alternative which are reported in `likfit`  
418 output.

### 419 **3.3. Validation**

420 Validation of each selected model or random forest was done using the validation data set. At  
421 each validation location the predicted soil pH was computed, and the prediction error was  
422 calculated as the difference between the predicted and observed soil pH (so a positive error is  
423 when the predicted pH exceeds the observed value). As exploratory summary statistics the  
424 mean error, median and mean square error were computed.

425 The validation data belong to a subset of SEA from the original survey, and as such are strongly  
426 clustered. Because of this the sample average of the squared errors may not be a good estimate  
427 of the mean square error, because the observations are not independent. A model-based  
428 approach was therefore taken to compute the expected squared error of prediction. A LMM  
429 was fitted to the prediction errors at the validation site (with a constant mean the only fixed  
430 effect). The expected square error (ESE) for each set of predictions was then computed as the  
431 sum of the squared mean error and the variances (nugget and spatially correlated) from the  
432 LMM. This is the *a priori* mean square error, i.e. the expected square error at a random location,  
433 and as such is likely to exceed the MSE computed directly from the errors of clustered data.

## 434 **4. RESULTS**

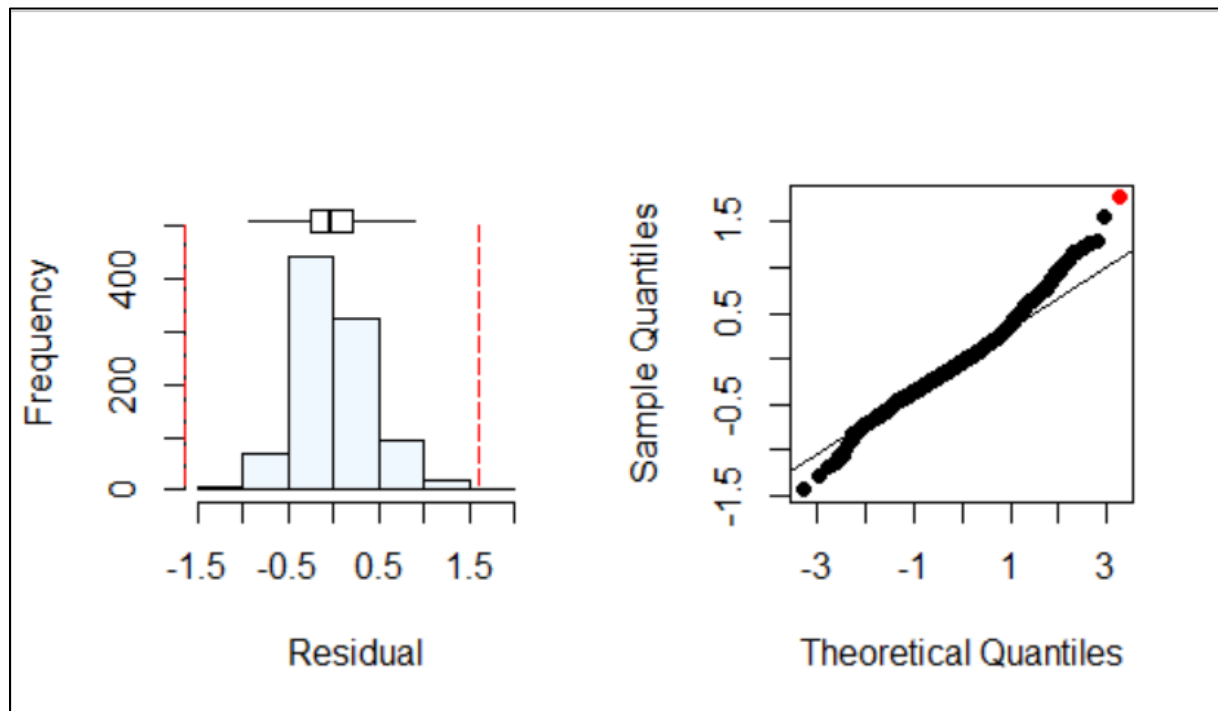
### 435 **4.1 Variable selection**

436 Table 2 and Figure 3 show the distribution of the residuals from the exploratory model. The  
437 histogram appears symmetrical and normal and the points on the QQ plot are close to a straight  
438 line. The residuals have octile skewness inside the range  $[-0.2, 0.2]$  and skewness inside  $[-1, 1]$ ,  
439 which would mean that a transformation is not normally considered necessary (Rawlins et al.,  
440 2005, Webster and Oliver, 2007).

441 *Table 2: Statistical summary of residuals from the exploratory model*

mean	Median	Variance	SD	Skewness	Octile skewness	Kurtosis
2.366e-16	-0.040	0.167	0.408	0.487	0.159	1.044

442



443

444 *Figure 3: Histogram and quantile plot of residuals from exploratory model*

445

446 *Table 3: REML estimates of parameters and AIC for the exploratory model, null model and the*

447 *hypothesis tests.*

Test	Predictors	Partial Sill	Range	Nugget	AIC	
					Max.likelihood	Non-spatial
	Exploratory (all predictors)	0.1367	21.08	0.218	1570	1664
0	mean	0.259	68.28	0.224	1618	1946
1	Rainfall	0.217	51.87	0.222	1611	1865
2	Rainfall + Soilclass	0.197	52.71	0.221	1622	1817
3	Rainfall + Slope	0.214	51.98	0.222	1610	1856
4	Rainfall + Elevation	0.180	39.70	0.220	1598	1774
5	Rainfall + Elevation + Valley depth	0.173	33.48	0.218	1598	1770
6	Rainfall+ Elevation + Landcover	0.174	40.09	0.220	1603	1769
7	Rainfall + Elevation + NDVI	0.181	39.83	0.220	1600	1776
8	Rainfall + Elevation + LS	0.174	38.84	0.221	1596	1760

9	Rainfall + Elevation + LS + CNBL	0.171	37.40	0.221	1597	1783
10	Rainfall + Elevation + LS + RSP	0.171	38.13	0.221	1598	1751
11	Rainfall + Elevation + LS + Aspect	0.173	39.18	0.221	1598	1760

448

449 Table 4 shows the log likelihood ratio and p-values of each test at respective degree of freedom  
450 df and chi-square distribution values. The likelihood ratios of tests 1,4 and 8 were greater than  
451 the chi-square distribution value values. Therefore, the null hypothesis for these cases were  
452 rejected, and the predictors retained during the sequential testing. The rest of the tests had log  
453 likelihood ratio less than their respective chi-square distribution values. Hence the null  
454 hypothesis was accepted for these predictors and they were dropped.

455 *Table 4: likelihood ratio and p-values of each hypothesis test at respective degree of*  
456 *freedom(df) and chi-square distribution values.*

Test		df	Chi-square	Likelihood ratio	p-value
1	Rainfall	1	3.841	9.533	0.002
2	Rainfall + Soilclass	17	27.587	22.506	0.166
3	Rainfall + Slope	1	3.841	2.429	0.119
4	Rainfall + Elevation	1	3.841	14.416	0.000
5	Rainfall + Elevation + Valley depth	1	3.841	1.817	0.177
6	Rainfall+ Elevation + Landcover	8	15.507	10.836	0.211
7	Rainfall + Elevation + NDVI	1	3.841	0.598	0.439
8	Rainfall + Elevation + LS	1	3.841	3.946	0.047
9	Rainfall + Elevation + LS + CNBL	1	3.841	1.047	0.306
10	Rainfall + Elevation + LS + RSP	1	3.841	0.403	0.525
11	Rainfall + Elevation + LS + Aspect	1	3.841	0.213	0.644

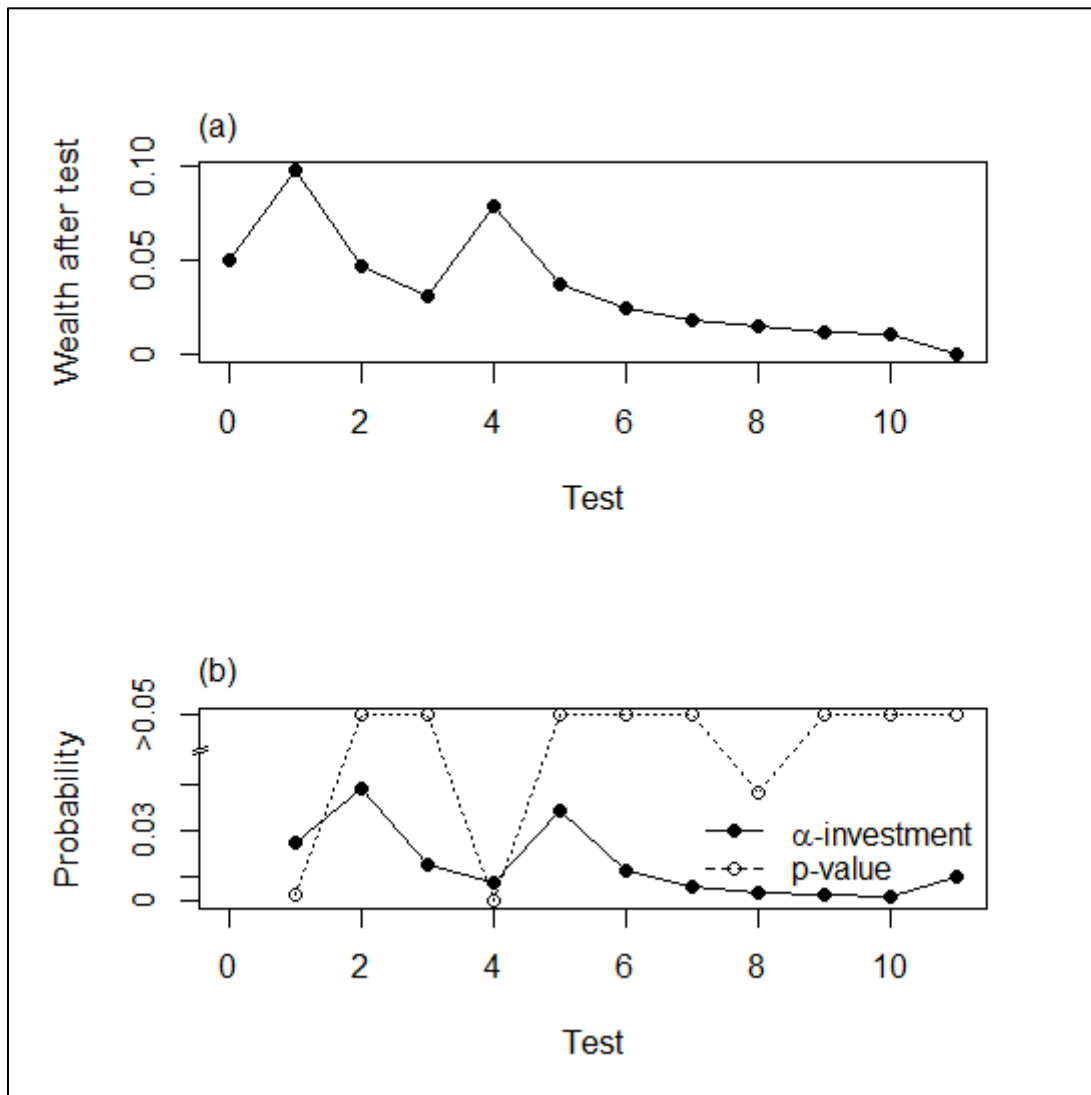
457 *Df= Degree of freedom which is the difference between the total degree of freedom for the*  
458 *target model and that of the model being compared to, LS= LS-Factor, CNBL= Channel*  
459 *Network Base Level, RSP= relative slope position.*

460

461 Figure 4.a shows the alpha wealth after each test and it can be observed that the quantity of the  
462 wealth is increased when the null hypothesis is rejected and depleted when the null hypothesis

463 retained, and it goes to zero at the end of the sequence. Figure 4.b shows the p-values (open  
 464 symbols) for the successive tests of additional predictors, as in Table 4, and the threshold (solid  
 465 symbols) against which each successive p-value is tested to achieve FDR. On this basis rainfall  
 466 and elevation were selected as predictors.

467



468

469 *Figure 4: a. alpha wealth after each test. b. probability of alpha investment and p-values*

470

471 **4.2 Spatial prediction of soil pH**

472 The estimated covariance parameters for the linear mixed models to be used for spatial  
 473 prediction of soil pH by the E-BLUP with elevation and rainfall as fixed effects for prediction  
 474 (Method A) and a constant mean as the only fixed effect (Method B, equivalent to ordinary  
 475 kriging) are shown in Table 5. The nugget, partial sill and range for the model with rainfall and  
 476 elevation as fixed effects are 0.220, 0.195 and 33.95 respectively. These values are smaller  
 477 than the corresponding parameters for the model with a constant mean as the only fixed effect  
 478 (0.224, 0.269 and 72.83). AIC values for both models are less than those of respective non-  
 479 spatial AIC. On this basis we may conclude that there is evidence for spatial dependence in the  
 480 random component of the LMM, and so potentially benefits in computing the E-BLUP for  
 481 spatial prediction at unsampled sites.

482 *Table 5: Covariance parameters for A= REML-EBLUP with elevation and rainfall as fixed*  
 483 *effects selected through alpha-investment (kriging with external drift), B=REML-EBLUP with*  
 484 *the only fixed effect a constant mean (ordinary kriging).*

Method	Partial Sill	Range	Nugget	AIC	
				Max.likelihood	Non-spatial
A	0.195	33.95	0.220	1184	1310
B	0.269	72.83	0.224	1614	1945

485

486 Table 6 shows the number of trees, number of predictors, number of variables considered at  
 487 each split, target node size and out-of-bag cross validation of the two random forest methods.  
 488 The out-of-bag MSE and R-squared show that there is a slight reduction in performance of the  
 489 random forest model with rainfall and elevation

490 *Table 6: ntree = number of trees in the forest; mtry = number of variables considered at each*  
 491 *split*

Method	ntree	predictors	mtry	Target node size	Out-of-Bag MSE	Out-of-Bag R-squared
--------	-------	------------	------	------------------	----------------	----------------------



C	200	11	3	5	0.31	0.30
D	200	2	1	5	0.32	0.27

492

493 Table 7 shows the permutation variable importance for each predictor when a random forest  
 494 model is fit with all predictors alone and when we include null predictors (sim1 to sim6) which  
 495 were generated by simulation to examine how random forest variable importance performs with  
 496 predictors that have no relation to the data. For the random forest model, the most important  
 497 variable is elevation with importance value of 0.166, followed by Channel Network Base Level  
 498 with value of 0.155. some variable importance values for some predictors are almost equal or  
 499 even less, but their p-values are much smaller. Null variables sim1 and sim6 despite have low  
 500 variable importance values, but very small p-values (less than 0.01). The inclusion of these  
 501 null variables has a substantial effect on the p-values of some predictors such as soilclass, slope,  
 502 landcover.

503 *Table 7: Permutation variable importance and p-values when a random forest model is fit with*  
 504 *all predictors alone and when we include null predictors (sim1 to sim6).*

Predictor	No null predictors		Null predictors included	
	Importance	p-value	Importance	p-value
rain	0.0889	0.0099	0.0912	0.0099
soilclass	0.0231	0.0198	0.0153	0.0099
slope	0.0318	0.8218	0.0268	0.2079
elevation	0.1657	0.0099	0.1718	0.0099
valley	0.0761	0.0099	0.0544	0.0099
landcover	0.0063	0.4752	0.0074	0.0792
NDVI	0.0499	0.0099	0.0470	0.0099
ls	0.0467	0.3267	0.0286	0.1287
cnbl	0.1554	0.0099	0.1477	0.0099
rsp	0.0554	0.0198	0.0330	0.0495
aspect	0.0148	0.1188	0.0066	0.3663
Sim1			0.0279	0.0099
Sim2			0.0232	0.0198
Sim3			0.0187	0.0594
Sim4			0.0204	0.0198
Sim5			0.0160	0.1782
Sim6			0.0274	0.0099

505

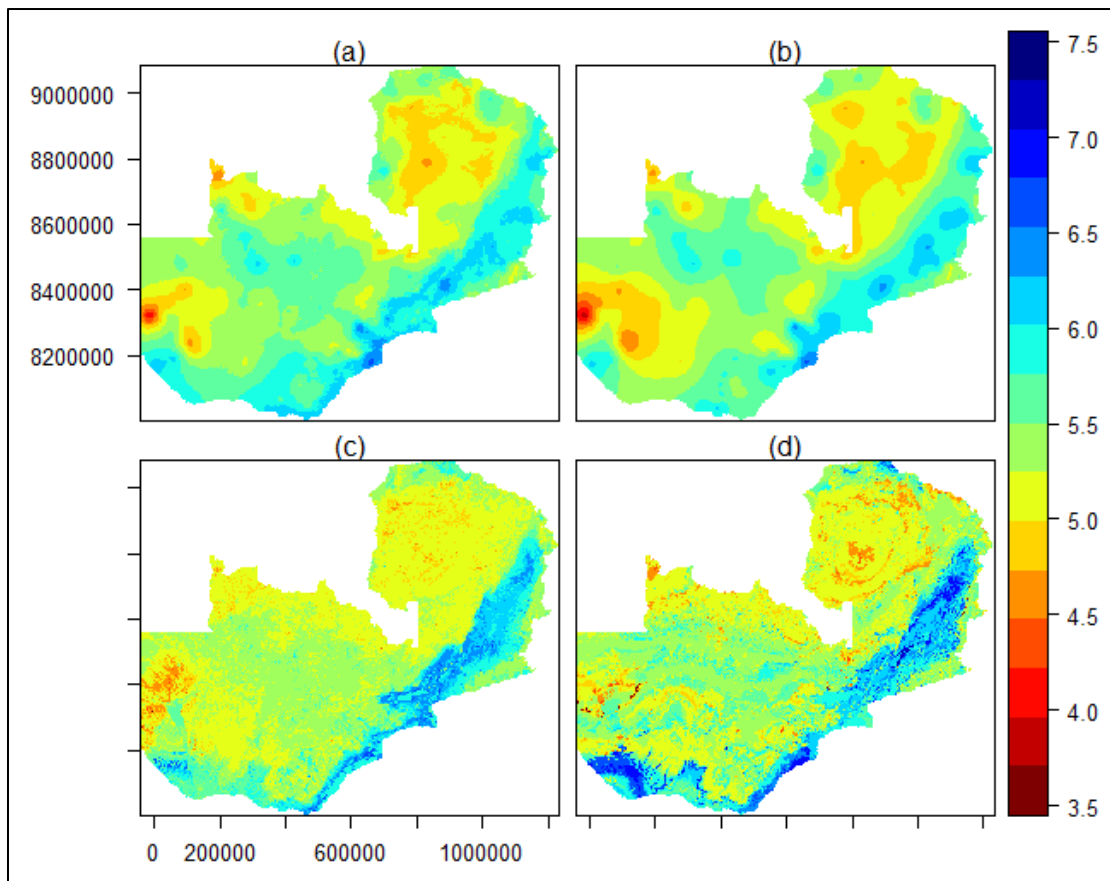
506 Table 8 shows the estimated parameters of the random forest residuals for the exponential,  
507 spherical and pure nugget correlation models. the non-spatial model was preferred because the  
508 AIC values for the spatial component was higher than that of the non-spatial component in both  
509 random forest predictions. Indeed, the fitted correlated variance for the spatial covariance  
510 function was zero. On this basis there is no scope to improve the RF predictions by a kriging  
511 step.

512 *Table 8: Estimated parameters of the exponential, spherical and pure nugget correlation*  
513 *functions for the residuals of the two random forest predictions.*

Method	Parameter	Exponential	Spherical	Pure.nugget
<b>RF (dem + rain)</b>	Partial Sil	0	0	0.188
	range	0	0	50
	Nugget	0.188	0.188	0
	AIC <sub>max.likelihood</sub>	1123	1123	1123
	AIC <sub>non-spatial</sub>	1119	1119	1119
<b>RF (all predictors)</b>	Partial Sil	0	0	0.157
	range	0	0	50
	Nugget	0.157	0.157	0
	AIC <sub>max.likelihood</sub>	951.4	951.4	951.5
	AIC <sub>non-spatial</sub>	947.4	947.4	947.4

514

515 Figure 5 shows the predicted spatial variability of soil pH. The spatial pattern is similar for all  
516 the models with low pH values (less than 5.5) in the Western and Northern parts and higher  
517 values (above 6) in the Southern and Eastern parts.



518

519 *Figure 5: Prediction maps of soil pH for (a) REML-EBLUP with elevation and rainfall as fixed*  
 520 *effects selected through alpha-investment (kriging with external drift), (b)REML-EBLUP with*  
 521 *the only fixed effect a constant mean (ordinary kriging (c) random forest with all predictors*  
 522 *(d). random forest with elevation and rainfall as predictors selected through alpha-investment.*

### 523 **4.3 Map Validation**

524 Table 9 shows the summary validation statistics for (A) REML-EBLUP with elevation and  
 525 rainfall as fixed effects selected through alpha-investment (kriging with external drift),  
 526 (B)REML-EBLUP with the only fixed effect a constant mean (ordinary kriging), (C) random  
 527 forest with all predictors (D). random forest with elevation and rainfall as predictors selected  
 528 through alpha-investment. The mean and median error values were smallest for the REML-  
 529 EBLUP (ordinary kriging) method while that of the REML-EBLUP (kriging with external drift)  
 530 was larger than that of the two random forest methods. The MSE and RMSE for the two REML-

531 EBLUP methods were smaller than those of the two random forest methods with REML-  
 532 EBLUP (ordinary kriging) having the smallest values. There was spatial dependency in the  
 533 prediction error in all the cases with the two REML-EBLUP cases having the smaller partial  
 534 sill values of 0.15 compared to 0.30 and 0.22 values for Random forest (elevation and rainfall  
 535 predictors) and random forest (all predictors) respectively. The ESE values for all the cases  
 536 were larger than the MSE values because the bias (ME) for the models is greater than zero.  
 537 REML-EBLUP (ordinary kriging) had the smallest ESE value.

538 *Table 9: Summary Validation statistics for(A) REML-EBLUP with elevation and rainfall as*  
 539 *fixed effects selected through alpha-investment (kriging with external drift), (B)REML-EBLUP*  
 540 *with the only fixed effect a constant mean (ordinary kriging), (C) random forest with all*  
 541 *predictors (D). random forest with elevation and rainfall as predictors selected through alpha-*  
 542 *investment.*

Variable		A	B	C	D
Prediction error	Mean	0.168	0.094	0.116	0.128
	Median	0.212	0.148	0.200	0.200
MSE		0.417	0.388	0.463	0.551
Corr.Model		Exponential	Exponential	Exponential	Exponential
Partial Sil		0.154	0.145	0.218	0.299
Range		48.380	44.670	40.200	36.250
Nugget		0.240	0.240	0.237	0.238
ESE		0.422	0.393	0.468	0.553

543

## 544 **5. DISCUSSION**

545 The mapped soil pH by all approaches is shown in Figure 5. The optimum pH (CaCl<sub>2</sub>) for plant  
 546 growth is between 5.2 – 7.5. bellow the pH of 5.2, the levels of Aluminum, Manganese and  
 547 Copper are toxic for plant growth, Phosphorous and Magnesium are not available to plant.  
 548 Above pH of 7.5, the interactions between Calcium, Magnesium and Potassium have a negative  
 549 impact on root absorption. Copper, Iron, Manganese, Zinc, Boron and Phosphorous are

550 deficient (Lake, 2000). The maps in Figure 5 show pH values less than 5.2 in the western and  
551 northern parts of the country meaning we expect these areas to have challenges of Aluminum,  
552 manganese and copper toxicity as well as Phosphorous and Magnesium deficiencies. In the  
553 Southern parts of the country the pH values, according to all the maps in Figure 5, range from  
554 5.2 to 7.5 which are optimal for plant growth. There are few areas in the southern part of Zambia  
555 with pH above 7.5. Similar spatial variations were observed by Chapoto et al., 2016. The  
556 southern parts where the pH is high is a valley area, the northern parts receives high rainfall and  
557 the western part despite receiving the same amount of rainfall as the eastern part, the areas is  
558 characterized by Kalahari sand. Our results show a similar spatial pattern for soil pH as that  
559 presented in the SoilGrids map ([www.soilgrids.org](http://www.soilgrids.org)) of Hengl et al., (2017). The main  
560 difference is that our map shows low pH values in the west of the country, whereas the SoilGrids  
561 map shows larger values there. Our results are more plausible pedologically given the parent  
562 material, and it has been long-established that the soils formed over the Kalahari sands of  
563 western Zambia are weakly to extremely acidic (Brammer, 1976). A more thorough assessment  
564 of the SoilGrids predictions using the RALS data would be of interest.

565 Predictions by the E-BLUP from the LMM with the only fixed effect a constant mean  
566 (equivalent to ordinary kriging) were better than other predictions in the sense that the mean  
567 and median errors were closest to zero and the mean square error and expected square error  
568 were the smallest. This is unexpected, given the evidence provided in the model-fitting stage  
569 for a significant relationship between soil pH and the selected covariates. This might be  
570 expected to result in better predictions from the LMM which includes these covariates as fixed  
571 effects. However, one may note (Table 5) that the correlated random variance in the LMM with  
572 rainfall and elevation as fixed effects is only about 25% smaller than the corresponding variance  
573 in the LMM with a constant mean the only fixed effect. The fact that a covariate is significantly  
574 related to a soil property does not necessarily mean that it will allow improved prediction of

575 that property relative to a model without that covariate. That is because the corresponding fixed  
576 effect coefficient must be estimated, and this estimation is a source of error in the prediction.  
577 Furthermore, Zimmerman et al., (1999) found that ordinary kriging performed better than  
578 universal kriging (UK, equivalent to the E-BLUP with some covariates) with a spatially  
579 clustered data set, while UK performed better when the data were not clustered. This may be  
580 because, in a strongly clustered data set, the effective degrees of freedom with which the fixed  
581 effects coefficients are estimated may be relatively small.

582 The use of random forests to include the environmental covariates in spatial prediction was less  
583 successful than the LMM and E-BLUP, with larger values of ESE. This could be due to over-  
584 fitting. It is notable that the residuals from the fitted RF at the calibration data points showed  
585 no spatial dependence, while the RF prediction errors at the validation points (Table 9) do show  
586 spatial dependence. This could arise because the RF algorithm, given its flexibility and ability  
587 to fit non-linear relationships, generates a model which closely fits the variations within the  
588 training data set, but which is not representative of the relationship between the predictor  
589 variables and target variable in the underlying population. This would lead both to marked bias  
590 in models of the random variation based on the residuals, as can also occur with ordinary least  
591 squares (Lark et al., 2006) and also in poor performance of the RF on a separate validation data  
592 set. These data may also provide a problem for the RF methodology because of the strong  
593 spatial clustering. If some data from a cluster are used in the development of trees while others  
594 are in the OOB subset then the assessment of the model and the value of the predictors may be  
595 over-optimistic. A predictor variable overfitted to a clustered data set might well fail to predict  
596 effectively at independent validation points. This emphasizes the importance of a genuinely  
597 independent validation of spatial predictions (Brus et al., 2011).

598 Spatial clustering of the observations may also be a contributing factor to the small p-values  
599 attributed to the entirely random, although spatially autocorrelated, null predictors which we

600 evaluated. This gives reason for caution when interpreting RF output. It is consistent with the  
601 findings of Wadoux et al., (2020) that the RF algorithm may select as predictor variables  
602 covariates which are not related to the target properties of interest by any direct or indirect  
603 causal relations. A spatially dependent predictor variable of this nature may indeed support  
604 spatial prediction of a variable to which it has no underlying relationship, but if this is the case  
605 then one might prefer to use a properly-designed set of orthogonal polynomial basis functions  
606 for the model rather than arbitrary variables. Furthermore, with strongly spatially clustered  
607 data, it is even more likely that a uncausally-related predictor will result in poor predictions at  
608 independent validation sites.

609 Many digital soil mapping studies use legacy data sets, rather than new samples collected for  
610 the purpose. As legacy data sets may originate in local surveys, or from networks of  
611 experimental stations, they may show marked spatial clustering, as do the RALS data because  
612 of their two-stage cluster sampling design. We note that such clustering may cause difficulties  
613 for the RF algorithm but that it is also important to account for it when dividing data into  
614 prediction and validation subsets. There is a risk of bias in the validation of a map if validation  
615 and training data are drawn from common clusters. The estimation of validation statistics from  
616 a validation set which is strongly clustered may also result in bias, which is why we have used  
617 a model-based approach to compute these statistics in this study. The expected squared error,  
618 computed from the model, in this case is not very different from the mean squared error  
619 computed directly although in each case it exceed the mean square error as expected (Table 9).  
620 This could be because the clusters are reasonably balanced (similar numbers of observations in  
621 each), and are themselves selected independently and at random. The model-based method to  
622 quantify prediction uncertainty is, nonetheless, a more general approach for use with validation  
623 data from locations not selected by probability sampling.

624 **6. CONCLUSION**

625 Spatial variability of soil pH was mapped using REML-EBLUP with elevation and rainfall as  
626 fixed effects selected through alpha-investment (kriging with external drift), REML-EBLUP  
627 with the only fixed effect a constant mean (ordinary kriging), random forest with all predictors  
628 and random forest with elevation and rainfall as predictors selected through alpha-investment  
629 models. The soil pH maps from these models showed similar patterns with pH values less than  
630 5.2 in the Western and Northern parts of the country. In the Southern parts of the country the  
631 pH values range from 5.2 to 7.5 which are optimal levels of soil pH for plant growth.

632 The ME, MSE and ESE (computed as the sum of the squared mean error and the nugget and  
633 spatially correlated variances from the LMM) were used to validate the performance of the  
634 models for spatial prediction of soil pH. The values of the ME, MSE and ESE from the  
635 validation statistics showed that REML-EBLUP with the only fixed effect a constant mean  
636 (ordinary kriging) performed better than the other methods.

637 Random forests had the largest values of MSE and ESE. This may result from over-fitting, and  
638 from the strongly spatially clustered distribution of the observations in the legacy data set which  
639 could affect the internal cross-validation in the RF algorithm.

640 We also noticed that the algorithm appeared susceptible to wholly random “null predictors”  
641 which we had simulated. Other studies, notably by Wadoux et al. (2020) have shown this, but  
642 we believe this to be the first example where mutually independent random spatially  
643 autocorrelated candidate predictor variables have been selected alongside pedologically  
644 plausible ones. The selection of such null predictors should give pause as it suggests that the  
645 random forest algorithm may be prone to overfitting. We suggest that this problem warrants  
646 further study as pedometricians should always aim to generate insight from their analyses, as  
647 well as predictions.



648 We note that legacy data, often used in digital soil mapping, may be strongly clustered like  
649 ours. We emphasize again the importance of splitting data into prediction and validation  
650 subsets at cluster level (i.e. allocating all data in any one cluster to the prediction or to the  
651 validation set). In this case there was not a large difference between the ESE (model-based  
652 estimate of the expected squared error) and the MSE (average of the squared errors), but this  
653 would not be true in general, and the use of a model-based approach to the analysis of validation  
654 errors at locations not selected independently and at random is most appropriate.

655 Finally, we note that we found no evidence for spatial correlation in the residuals from the fit of the  
656 random forest to the prediction data set, although there was correlation in the prediction errors by this  
657 method at the validation sites. This is an important reminder that such residuals given us little if any  
658 insight into the actual behavior of the prediction error, and are good reasons to avoid using kriging  
659 methods in combination with modelling methods which, unlike REML and EBLUP, do not have a built-  
660 in methodology to estimate parameters of the error without bias.

## 661 **7. ACKNOWLEDGEMENT**

662 This study was made possible with support from the Zambian Ministry of Higher Education  
663 through their 2016/2017 science and technology female postgraduate scholarships; and from  
664 the Commonwealth Scholarships Commission through their split-site scholarship funded by the  
665 UK government. We also thank Indaba Agricultural Policy Research Institute (IAPRI) for  
666 allowing us to use their RALS 2012 survey data.

667

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