

Abstract

 Arsenic (As) in native soils of the Amazon rainforest is a concern due to its likely origin from the Andean rivers, which transport loads of sediments containing substantial amounts of trace elements coming from the cordilleras. Yet, unveiling soil As baseline concentrations in the Amazon basin is still a need because most studies in Brazil have been performed in areas with predominantly high concentrations and cannot express a real baseline value for the region. In this study, 414 soil samples (0-20, 20-40 and 40-60 cm layers) were collected from different sites throughout the Amazon basin - including native Amazon rainforest and minimally disturbed areas - and used to determine total and extractable (soluble + available) As concentrations along with relevant soil physicochemical properties. Descriptive statistics of the data was performed and Pearson correlation supported by a Principal Component Analysis (PCA) provided an improved understanding of where and how As concentrations are influenced by soil attributes. Total As concentration ranged from 0.98 to 41.71 mg kg^{-1} with values usually increasing from the topsoil (0-20 cm) to the deepest layer (40-60 cm) in all sites studied. Considering the proportional contribution given by each fraction (soluble and available) 37 on extractable As concentration, it is noticeable that KH₂PO₄-extractable As represents the most important fraction, with more than 70% of the As extracted on average in all the sites studied. Still, the extractable fractions (soluble + available) correspond to ~0.24% of the total As, on average. Total, 40 available, and soluble As fractions were strongly and positively correlated with soil Al^{3+} . The PCA indicated that soil pH in combination with CEC might be the key factors controlling soil As concentrations and the occurrence of each arsenic fraction in the soil layers.

Keywords: arsenic; extractable As; prevention level; carcinogenic; baseline; Amazon.

1. INTRODUCTION

51 in 15 of the 20 Latin American countries in 2011, and much of the As contamination described during

52 the last 20 years in this region were published much later or never in the international literature

53 (Bundschuh et al., 2012).

The overall mean value of total As for different soils worldwide is estimated as 6.83 mg kg⁻¹ 55 (Kabata-Pendias, 2011), which is very close to the mean value found for Brazilian oxisols (i.e., 5.92 56 mg kg⁻¹) (Campos et al., 2007). Yet, As is found more abundantly in argillaceous sediments (up to 13 57 \parallel mg kg⁻¹)-than in any other rock. Anionic forms, AsO₂, HAsO₄², and H₂AsO₃², predominate in soil 58 solution (Kabata-Pendias and Mukherjee, 2007).

 Exposure to As has a great impact on human health across the world (Kapaj et al., 2006; Meliker et al., 2007; Chatterjee et al., 2018). In the environment As often raises a great concern because of its high toxicity to diverse human body systems (e.g. integumentary, nervous, respiratory, cardiovascular, hematopoietic, immune, endocrine, hepatic, renal, reproductive) as well as to its 63 capability of causing different types of cancer (skin, lung, and bladder) (Mohammed Abdul et al., 2015; Gamboa-Loira et al., 2017). In fact, this naturally occurring metalloid is considered a well- established "Class I" human carcinogen by the International Agency of Research on Cancer (IARC) (Minatel et al., 2018).

 Arsenic contamination originating from geogenic sources occurs in many regions worldwide, such as Bangladesh and India (Arco-Lázaro et al., 2018; Rahman et al., 2018). It is estimated that over 200 million people are chronically exposed to As concentrations at or above the World Health Organization (WHO) threshold across the world (Naujokas et al., 2013). In addition to the natural 71 Sources occurrence of arsenic, there are many other anthropogenic sources that may increase soil arsenic concentration, for example: mining, smelting, application of wastes, animal manures and As- bearing pesticides, wood preservation, irrigation using water contaminated, coal fly ash, glass manufacturing, wastewater sludge, pharmaceutical waste, livestock dips, smelting activities to phosphate fertilizers (Punshon et al., 2017).

76 Several studies have clearly shown that the adverse effects of arsenic are not only dependent 77 on the total concentration of this element, but rather on its reactive fractions (Wang et al., 2018). For 78 As, this fractionation may include soluble and exchangeable, bound to carbonates, iron and manganese

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 oxides, organic matter, sulphides and residual fraction among others. Those fractions govern arsenic bioavailability and mobility in soils (Ho et al., 2013; Shahmoradi et al., 2017; Liu et al., 2018). Additionally, soil physicochemical properties such as pH, redox potential, clay and organic matter (OM) content and the presence of metal (Fe, Mn and/or Al) oxides and/or competing ions such as 83 phosphates highly influence arsenic bioavailability and mobility (Sharma et al., 2010; Biswas et al. 84 2014).

 Many sequential extraction procedures (SEP) can be used to extract chemical fractions of trace elements in soil (Wan et al., 2017; He et al., 2018). These methods classify trace elements in soil into different operationally-defined fractions with increasing metal binding strength (Wan et al., 2017). Soil chemical extractions for soluble and available As can help understand the element's dynamic chemical behavior and its possible entrance into the food chain.

90 Many extractants can be used for this purpose, for example, the soluble fraction can be 91 extracted using water (water-soluble) (Wenzel et al., 2001; Baroni et al., 2004; Jiang et al., 2014; 92 Wan et al., 2017; Baroni et al., 2004; Jiang et al., 2014). This fraction contains water-soluble species 93 such as free arsenate, and weakly adsorbed species, representing the most mobile and potentially the 94 most-bioavailable As species in the environment (Wan et al., 2017). Additionally, the soluble fraction 95 can be also extracted using inorganic salts such as KCl (Kim et al., 2003) and KNO₃ (Alam et al., 96 2007).

97 Another relevant fraction of As - the available fraction - is not well defined in the literature in 98 terms of method of extraction, but would be better classified as the As retained in the solid phase 99 interacting with the soil solution and occupying exchangeable sites as difuse ion or as outer-sphere 100 complexes (Van Herreweghe et al., 2003) and can be extracted by using Na₂HPO₄ (Kim et al., 2003), 101 MgCl₂ (Jiang et al., 2014), CH₃COOH (Wan et al., 2017), $(NH_4)_2SO_4$ (Wenzel et al., 2001) and 102 KH₂PO₄ (Thinh et al., 2018) among others.

 In Brazil, soil quality criteria for some contaminants were defined by CONAMA in Normative 420/2009, which sets screening values based on total concentration of selected elements. For arsenic, 105 the established threshold for a so-called "prevention level" in soil was 15 mg As kg^{-1} and is primarily based on phytotoxicity tests and ecological risk assessment. Besides this, investigation/intervention **Field Code Changed Formatted:** Font: 11 pt

107 values derived from human health risk assessments were set as 35 mg As kg^{-1} for agricultural, 55 mg 108 As kg⁻¹ for residential, and 150 mg As kg⁻¹ for industrial areas (Conama, 2009). These thresholds were established for total concentrations and in some situations can be very restrictive, especially in soils with low bioavailability. This justifies investing more effort to assess As availability in soils, especially in tropical ecosystems, which are known for their oxidic characteristics that render arsenic less bioavailable (Campos et al., 2007; 2013a).

 Increased As levels have been observed in soils worldwide and also in Brazil, especially due to mining activities (Bundschuh et al., 2012). Hypothetically, we could expect high concentrations of As in soils nearby the Amazon basin because of the sediments coming through the Amazon river (and its tributaries) from the Andes, which are expected to be naturally enriched in selected volatile trace elements (e.g., Se and Hg) (Fadini and Jardim, 2001; Silva Junior et al., 2017). However, there is a lack of comprehensive information about background levels of As in soils from the Amazon basin to prove this assumption since the more recent studies have focused only in specific states of Eastern Amazon (Fernandes et al. 2018). Therefore it is necessary to perform a more inclusive study concerning the occurrence/distribution of As in different Amazon agroecosystems in order to provide information for the local population about possible health issues involved in the use of soil and water resources for living, building and growing crops in this region.

 To the best of our knowledge, no research data regarding the determination of soluble and available soil As in the Brazilian Amazon region have been published so far. This is relevant, as such information could support public policies for decision makers and help guide future research in these sites in order to provide a better planning for a sustainable use of land and natural resources by the local population, while ensuring also human health and food security.

 In this context, this study aims to: i) characterize and compare baseline levels of As in different sites in the Amazon region; ii) study the dynamics of extractable As fractions (soluble and available fractions) with depth, i.e., in three different layers (0-20, 20-40 and 40-60 cm) of soils occurring in agroecosystems with Brazil nuts plantations; and, iii) assess the influence of the main soil physicochemical attributes for determining the soil As concentrations and the occurrence of each arsenic fraction in the soils.

2. MATERIALS AND METHODS

2.1 Study approach

 This study involved the following steps: (a) soil survey to collect baseline data for As in 9 target sites; (b) description of the natural variation of As and soil physicochemical attributes in each site; and, (c) correlation and PCA analysis to support inference of soil environmental controls of As variability in the soil. The methods used to reach our goals are described in the following sections.

2.2 Sampling sites and characterization of the study area

 Soil samples were collected in nine sites belonging to the Amazon Biome in Brazil. Each site correspond to a Brazil nut stand, identified by: SM (Sena Madureira, state of Acre); XP (Xapuri, state of Acre); ANO (Anori, state of Amazonas); AR (Aruanã farm-Itacoatiara, state of Amazonas); AP (Laranjal do Jari, state of Amapá); MT (Itaúba, state of Mato Grosso); PA (Santarém, state of Pará); RO (Porto Velho, state of Rondônia) and RR (Caracaraí, state of Roraima). In each site, samples of 15 149 collection points were **collected** sampled in the native Amazon rainforest, except for AR (Itacoatiara), where samples of 18 spots were collected in the Aruanã farm, which is a large Brazil nuts plantation. Samples were collected between 2014 and 2017. The chosen target sites for this study were the same ones used by Silva Junior et al. (2017) in their survey assessing the natural variation of selenium in Brazil nuts and soils from the Amazon region, to which we have added four additional areas. Sampling sites and detailed geographic information are provided in **Figure 1 and Table S1**.

 Soil samples were collected at the depths of 0-20, 20-40, and 40-60 cm at a distance of 3 m from the trunk of selected Brazil nuts trees. These samples were collected at four points and mixed to compose a single sample for each depth. The sampling method is represented schematically according to Silva Junior et al. (2017).

161 **Figure.** 1. Map of the Brazilian Amazon (within South America) (A) with visualization of the Andes nountain range and Amazon sub-regions (B) as well as indicating the location of the nine studied site 162 mountain range and Amazon sub-regions (B) as well as indicating the location of the nine studied sites 163 | (C) identified by: Sena Madureira (SM) and Xapuri (XP)—state of Acre; Porto Velho (RO)-state of 163 (C) identified by: Sena Madureira (SM) and Xapuri (XP)—state of Acre; Porto Velho (RO)-state of 164 Rondônia: Anori (ANO) and Itacoatiara (AR)-state of Amazonas: Itaúba (MT)-state of Mato Grosso: 164 Rondônia; Anori (ANO) and Itacoatiara (AR)-state of Amazonas; Itaúba (MT)-state of Mato Grosso;
165 Caracaraí (RR)-state of Roraima; Santarém (PA)-state of Pará and Laranjal do Jari (AP)-state of 165 Caracaraí (RR)-state of Roraima; Santarém (PA)-state of Pará and Laranjal do Jari (AP)-state of 166 Amapá. Relevant locations referred to in the manuscript for data comparison (e.g., the Cerrado b 166 Amapá. Relevant locations referred to in the manuscript for data comparison (e.g., the Cerrado biome 167 as well as the State of Minas Gerais) are also shown in the map. DEM - Digital elevation model. as well as the State of Minas Gerais) are also shown in the map. DEM - Digital elevation model. 168 Source: U.S. Geological Survey's Center for Earth Resources Observation and Science (EROS) 169

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171 **2.3 Soil physicochemical characterization**

 Soil samples were taken to the Soil Science Department of the Federal University of Lavras, Minas Gerais-Brazil, where they were ground, sieved at <2 mm, dried at room temperature, and subjected to chemical and physical (texture) analysis. The soil pH was determined potentiometrically 175 in water (pH H₂O) and KCl (1 mol L⁻¹) in the ratio 1:2.5 soil/solution. The exchangeable cations (Ca²⁺, 176 Mg²⁺ and Al³⁺) were obtained by 1 mol L⁻¹ KCl extractants. Sulfur (S-SO₄²) was extracted as sulfate by monocalcium phosphate with acetic acid and the result converted to S. Available P, Na, K, and 178 micronutrients (Fe, Zn, Mn, and Cu) were obtained by Mehlich-1 (0.05 mol L^{-1} HCl + 0.0125 mol L^{-1} H₂SO₄; in ratio 1:10 soil/solution). Potential acidity $(H + Al)$ was estimated indirectly through SMP

180 pH (Shoemaker et al., 1961). Exchangeable aluminum $(AI³⁺)$ was determined by titration with sodium 181 hydroxide (0.025 mol L⁻¹ NaOH). P was determined colorimetrically, K by flame emission 182 photometry, S-SO₄² by turbidimetry, and Ca, Mg, Zn, Fe, Mn, and Cu were determined by atomic 183 absorption spectrometry (AAS) (Raij and Quaggio, 1983). The organic matter content was determined 184 by oxidation of the organic matter in the soil with a potassium dichromate solution in the presence of 185 sulfuric acid (Carter and Gregorich, 2006). Sum of bases (SB), cation exchange capacity at pH 7 (T), 186 effective cation exchange capacity (t), base saturation $(V\%)$, and aluminum saturation (m) were 187 calculated based on the results of the chemical analysis. The particle size analysis (clay, silt, and sand) 188 was performed by the "Bouyoucos" method, as described by Embrapa (1997).

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190 **2.4 Total arsenic determination**

191 Total As analyses was performed in air-dried and <2 mm sieved soil samples. A portion of 192 each soil sample was finely ground (< 0.45 mm) using an agate ball mill (Retch PM 400 mill). Soil 193 samples (~0.2 g DW) were fully digested in 70% HF, 70% HNO₃ and 60% HClO₄ (TAG; Fisher 194 Scientific UK Ltd), using perfluoroalkoxy (PFA) digestion vessels and a 48-place teflon-coated 195 graphite digester block (Model A3, Analysco Ltd, Chipping Norton, UK) (Chilimba et al., 2011).

 The digester block runs with two different programs, a mix of reagents and a few steps for 3 days as can be observed with further details in table 1. After the digestion, the block was turned off and the tubes were left to cool down completely for about 30 minutes. Digested samples were diluted 199 to 50 mL into volumetric flasks using Milli-Q water and stored in 5% HNO₃ at room temperature in universal sample bottles, pending elemental analysis. Before the analysis in the ICP-MS, the extracts 201 were diluted using $1 mL + 9 mL$ of Milli-Q water.

204 *Left overnight with the program turned on.

206 Each digestion batch (n = 70) included a standard reference material (n = 3) (NIST 2711a; 207 Montana Soil II, National Institute of Standards and Technology, Gaithersburg, MD, USA), which 208 contains 107 ± 5 mg kg⁻¹ As. The average recovery in the standard reference material was 101% (n = 209 21, SD = 6.61) for As. In addition, 3 blank samples were used in each batch for quality control 210 purposes and for calculating limits of detection. The limit of detection (LOD) was established using 21 211 blank extracts for soil samples following the overall procedure. The values were calculated with three 212 times the standard deviation (SD) of the blanks used in each batch $(n = 3)$ using the formula: SD (3) 213 blanks) x 3 x final volume extract (mL) ÷ sample weight (g). The average LOD was 0.029 mg kg⁻¹ for 214 total soil As.

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216 **2.5 Soluble and available As in soil (extractable As)**

 Soluble and available As analyses were performed in sequence, according to the modified method described by Cai et al. (2002). First, for the soluble fraction, 2 g of air-dried and <2 mm sieved 219 soil was added into a 50 mL polycarbonate centrifuge tube. Next 10 mL of 0.01 mol L^{-1} KNO₃ were added to each tube and the suspensions were shaken for 2 hours using a rotary shaker (20 rpm). Soil 221 suspensions were then centrifuged for 30 min at 3500 rpm, filtered to <0.22 µm using a Millex syringe driven filter unit (Millipore, Cork, Ireland) and stored at room temperature prior to analysis of As by 223 ICP-MS. In the second step, for the available fraction, an extraction with 10 mL of 0.016 M KH₂PO₄ was performed in the same tubes from the previous step. The suspensions were shaken for 1 hour in the rotary shaker (20 rpm) and then centrifuged for 30 minutes at 3500 rpm. The supernatant was filtered using <0.22 µm syringe filter and put in a universal tube (used for analysis in the ICP-MS). 227 For the ICP-MS analysis, 4.8 mL of the supernatant collected from KNO_3 and KH_2PO_4 extractions 228 were added into an ICP tube containing 0.2 mL of 50% HNO₃ for As analysis. The limits of detection (LOD) for soluble and available soil As were calculated as previously mentioned, and corresponded to 0.047 and 0.056 μ g kg⁻¹, respectively.

2.5 Analytical determination of As

 The analytical determinations of As were performed using Inductively Coupled Plasma-Mass Spectrometry (ICP-MS, X-Series II, Thermo Fisher Scientific) in the Environmental geochemistry 235 analytical suite from the University of Nottingham-UK. Samples were introduced at 1 mL min⁻¹ through a concentric glass venturi nebuliser and peltier-cooled (3 °C) spray chamber. Instrumental 237 drift was corrected using internal standards $(^{72}Ge, ^{103}Rh$ and ^{185}Re). Data for total As concentration in 238 soils are reported on a dry weight basis (DW) and expressed in mg kg^{-1} . Data for soluble and available 239 As are expressed in μ g kg⁻¹.

2.6 Statistical analysis

 A descriptive analysis of the concentrations of total As, as well as for soluble and available As was carried out using Box plots in order to display data distribution. The data were subjected to Pearson's correlation analysis, which was performed between total, soluble and available As concentrations and other physicochemical soil properties. All the graphics were performed using the software R 3.4.4 (R Development Core Team, 2018).

 Analysis of variance was performed to assess how the log-transformed As concentration in soils varied across different sites. After verifying the significance through ANOVA, the mean values for As concentration were compared among sites for the three layers studied using Tukey's HSD test with the package emmeans v 2.23 (Lenth, 2016) in R 3.4.4 (R Development Core Team, 2018). The same approach was used to compare soluble and available As in soils, but in this case using generalized least squares (GLS) for the data log-transformed, in order to allow the variance to be estimated separately for each site (Cleasby and Nakagawa, 2011).

 A principal component analysis (PCA) was performed to find out the main soil attributes influencing As concentration in the soil according to the degree of variance explained and to confirm the previous relationships obtained by Pearson correlation analyses. Because the distribution of As concentrations and fractions was not normal, the data were log-transformed, before analysis to meet the assumption of normality required for the Pearson correlation and PCA, except for soil pH, because it follows a normal distribution in soil data.

3 RESULTS AND DISCUSSION

3.1 Soil physicochemical properties, element contents and associations

 The chemical and textural attributes of soil varied widely among the sites studied, reflecting a great variability in soil characteristics throughout the Amazon basin that may be related to diversity of parent material and soil types. Most of collected samples show acidic soils (mean pH ranging from 3.75 to 5.5) which in fact reflects the well known weathered soils from this region according to Alvarez et al. (1999). The OM content in the topsoil is classified as moderate in most sites (average OM content ranging from 1.16% to 2.63%) according to Alvarez et al. (1999) (**Table 2**). The moderate contents of OM can be explained by the high temperature and precipitation levels (humidity) in the region, i.e., the most important climatic factors controlling the dynamic of ecological processes (da Silva et al., 2018), which reflects not only in the biomass production but also in decomposition rates of the litter in the soil.

 The dominant soil texture was clay (AR, RO and PA) and sandy loam (MT and AP). Such high percentage of clay content facilitates As retention, especially onto kaolinite, as well as Fe and Al oxyhydroxide surfaces that predominate in the clay fraction of the highly developed soils of the Amazon region, which is subjected to strong weathering conditions (de Souza et al., 2018) (**Table 2**).

277 Soil cation exchange capacity (CEC) ranged from 2.96 to 25.2 cmol_c dm⁻³ demonstrating a 278 great variability among the sites studied. The soils from ANO showed the highest mean values of CEC 279 (22.6 cmol_c dm⁻³). For the others sampled sites, CEC values were all below 13 cmol_c dm⁻³ with the 280 | lowest values observed in the sites RR and AP (5.33 and 4.28 cmol, $dm³$, respectively). This wide variation in soil CEC indicates that besides the predominance of dystrophic and highly weathered soils with secondary minerals, there are still soils with 2:1 clay minerals, composing less weathered soils with high activity in the clay fraction in this region (**Table 2**).

 Extractable S and P in the present study exhibited considerable variation among the sites with 285 mean values ranging from 5.29 to 47.9 mg kg⁻¹ (S) and from 0.52 to 3.72 mg kg⁻¹ (P). In general, the 286 content of available P was moderate in most sites, where also the OM contents are usually higher than in deeper layers, except in the sites MT and RO in southern Amazon, which contained available P

288 values below 2 mg kg⁻¹ (Table 2). The content of extractable S on the other hand presented higher values in the sites AR and RO (**Table 2**).

Site	n	pH(H ₂ O)	$OM(\%)$	CEC			Fe	Clay	Silt	Sand
				$(\text{cmol}_c \text{ dm}^{-3})$	$(mg kg-1)$	$P(mg kg-1)$	$(mg kg-1)$			
SM	15	4.53 ± 0.34	$.42 \pm 0.35$	6.92 ± 1.74	5.29 ± 1.17	3.72 ± 1.27	174.9 ± 86.3	21.6 ± 4.70	22.5 ± 2.38	55.9 ± 4.85
XP	15	4.36 ± 0.32	1.62 ± 0.55	7.21 ± 2.25	8.09 ± 5.16	3.08 ± 0.75	190.1 ± 83.9	18.4 ± 5.37	32.3 ± 10.2	49.3 ± 10.5
MT	15	4.19 ± 0.16	2.32 ± 0.40	7.10 ± 2.52	7.32 ± 2.00	1.46 ± 0.32	324.3 ± 69.6	19.5 ± 2.61	3.86 ± 1.85	76.6 ± 2.82
AR	18	4.26 ± 0.09	2.63 ± 0.32	7.97 ± 1.34	16.5 ± 7.06	2.56 ± 0.61	183.0 ± 48.6	77.6 ± 3.46	10.1 ± 2.37	12.3 ± 2.56
ANO	15	4.00 ± 0.35	2.03 ± 0.70	22.6 ± 10.9	9.55 ± 4.97	2.53 ± 0.85	489.7 ± 134.1	27.2 ± 5.82	56.7 ± 8.05	16.1 ± 11.9
AP	15	5.49 ± 0.34	1.16 ± 0.34	4.28 ± 0.67	9.49 ± 1.60	2.25 ± 0.72	58.5 ± 22.6	19.9 ± 2.60	3.40 ± 1.92	76.7 ± 3.24
RR	15	4.46 ± 0.16	2.43 ± 0.46	5.33 ± 0.73	10.1 ± 1.64	2.38 ± 0.28	177.2 ± 50.2	44.1 ± 5.97	8.13 ± 3.54	47.7 ± 6.12
RO.	15	3.75 ± 0.22	2.40 ± 0.37	12.6 ± 3.63	18.0 ± 8.88	1.87 ± 0.37	196.2 ± 70.8	51.7 ± 9.78	28.1 ± 12.5	20.2 ± 7.70
PA†	15	4.41 ± 0.26	2.20 ± 0.42	9.24 ± 2.24	10.6 ± 3.51	2.25 ± 0.43	242.4 ± 62.5	52.0	8.0 ₀	40.0

297 **Table 2.** Physicochemical properties of the topsoil (0-20 cm) from samples collected at different sites in the Amazon region.

Values of soil physicochemical attributes are mean \pm standard deviation (SD); †Texture in Santarém/PA was derived from one single analysis performed by a
composite sample resulted of subsamples from 5 spots in the site

composite sample resulted of subsamples from 5 spots in the site. Each site correspond to a Brazil nut stand, is identified by: SM (Sena Madureira, state of

300 Acre); XP (Xapuri, state of Acre); ANO (Anori, state of Amazonas); AR (Aruanã farm-Itacoatiara, state of Amazonas); AP (Laranjal do Jari, state of Amapá); 301 MT (Itaúba, state of Mato grosso); PA (Santarém, state of Pará); RO (Porto Velho, state of Rondônia) and RR (Caracaraí, state of Roraima).

3.2 Total As concentration in soil

303 Total As concentrations ranged from 0.98 to 41.7 mg kg⁻¹, with 75% of the soil samples 304 showing values lower than 6.93 mg $kg⁻¹$, which is close to the mean value of total As reported for | different soils worldwide (i.e., 6.83 mg kg⁻¹) (Kabata-Pendias, 2011). The sites with highest baseline 306 As concentration were AC:SM followed by RO and AC:XP. On the other hand, the sites with lowest As concentrations were AP and RR, according to the Tukey HSD test (**Figure. 2**).

308 Interestingly, total As concentration in the western Amazon (AC:SM, AC:XP and RO) presented the highest values and a wide variation not only among the layers but also within the three layers evaluated. The total As concentration in the western Amazon showed median concentrations 311 increasing from the topsoil to the deepest layers as follow: $\overline{AC:SM}$ (15.9 < 19.0 < 26.5 mg kg⁻¹) and 312 **AC:**XP (5.96 < 7.20 < 9.34 mg kg⁻¹) for the layers 0-20, 20-40 and 40-60 cm, respectively. A wide variation for arsenic concentration within layers was observed especially at the deepest layer (40-60 cm), which also presented the highest As concentrations for each site studied in the western Amazon: 315 | AC:SM (17.3 to 41.7 mg kg⁻¹), AC:XP (5.69 to 17.8 mg kg⁻¹), and RO (5.83 to 23.6 mg kg⁻¹) (**Figure**. **2**).

 There is no nationwide As baseline level established by the Brazilian legislation to be used in terms of quality reference values (QRVs) for soils. This is also true for soils of the Brazilian Amazon region. However, a QRV for As in soils from the Eastern Amazon was proposed by Fernandes et al. (2018) . These authors observed total As concentration ranging from 0.07 to 7.75 mg kg⁻¹ in the 0-20 321 cm layer and, considering the $90th$ percentile of the sample population after eliminating anomalies 322 (cumulative frequency distribution), 2.7 mg kg^{-1} was suggested as the QRV. Given the heterogeneity in the pedogeoclimatic conditions in the region, the $90th$ percentile seems to best represent the QRVs because of the wide variances for the values observed, as suggested by Conama (2009).

 In the present study 57% of the samples from the 0-20 cm soil layer show As concentrations higher than the QRV for the 90th percentile proposed by Fernandes et al. (2018), which reinforces the difficulty of establishing a single QRV for a very large region as the Brazilian Amazon. In addition, the medians for total As concentrations in the three layers (0-20, 20-40 and 40-60 cm) from the site | AC:SM are higher than the As prevention level (15 mg kg⁻¹) established by the Brazilian Normative CONAMA 420, for soils nationwide (Conama, 2009) (**Figure. 2**). This is indicative that the current Brazilian prevention values need to be validated using representative soils from the Amazon basin since this region may present soils that may be naturally enriched with As. All other studied sites 333 presented median As concentrations below 6 mg kg^{-1} - a value close to the average As content reported 334 by Campos et al. (2007) for 17 Brazilian oxisols (i.e., 5.92 mg kg⁻¹) - with the lowest median 335 concentrations observed in the site AP (1.30, 1.53 and 1.71 mg kg^{-1} for 0-20, 20-40 and 40-60 cm respectively) (**Figure. 2**).

337 The sites from the central and eastern portion of the Amazon basin (AM:ANO, AM:AR, PA and AP) presented low As concentrations in all the three layers evaluated. They also presented low variation among and within the layers, as observed in the medians concentrations in the site \overrightarrow{AM} :AR (5.07, 5.33 and 5.41 mg kg-1 for 0-20, 20-40 and 40-60 cm respectively) (**Figure. 2**).

 The sites located in the northern and southern portion of the Amazon, which are RR (medians 342 of 1.39, 1.70 and 1.80 mg kg^{-1}) and MT (medians of 2.56, 2.99 and 3.03 mg kg^{-1}) also presented low total As concentration for the three layers evaluated compared to the other sites studied, and no meaningful variation among and within the layers was noticed (**Figure. 2**).

 Campos et al. (2013b) determined background levels of As in representative soils from the states of Goiás and Minas Gerais (from 2000 to 2500 km far away from the studied sites), which belongs to Cerrado biome (Brazil) in three sub-regions and observed concentrations similar to those found in the southern Amazon in the present study, with mean concentrations in a decreasing order as 349 follow: East of state of Goiás $(3.29 \text{ mg kg}^{-1})$ > Triângulo Mineiro (state of Minas Gerais) $(2.18 \text{ mg kg}^{-1})$ ¹) > Northeast of Minas Gerais (0.62 mg kg⁻¹). The highest As concentrations were observed in profiles of Plintosols and Cambisols from the East of Goiás. In another study, Campos et al. (2013a) 352 observed levels of As ranging between 0.14 mg kg⁻¹ (for a gley soil) and 9.3 mg kg⁻¹ (for an oxisol) in soils from Minas Gerais. Those concentrations are usually found in non-contaminated soils (< 10 mg kg⁻¹) (Adriano, 2001; Campos et al., 2007, 2013a; Bundschuh et al., 2012) and are similar to the ones reported in the present study.

 In a different scenario far away from the Amazon region, in the Iron Quadrangle, southeast region of Brazil, within the state of Minas Gerais, the presence of As is associated with primary gold 358 deposits where it is contained predominantly in S-bearing minerals such as arsenopyrite and pyrite 359 (Borba et al., 2003). Also in the iron quadrangle, Alves and Rietzler (2015) observed high As 360 concentrations in soils from mining areas in Nova Lima (13.2 in the control site and 489 mg kg⁻¹ in the 361 mine stream) and Santa Bárbara (82.3 in the control site and 1329 mg kg⁻¹ nearby the mine plant). Out 362 of the Iron Quadrangle, but still in Minas Gerais, natural soil As concentrations found in a private 363 natural reserve (native Cerrado vegetation) located nearby a gold-mining area ranged from 25 to 35 364 mg kg⁻¹ (Ono et al., 2012), which is also greater than the prevention level established by Normative 365 420, for Brazilian soils (Conama, 2009). These studies emphasize the importance of the parent 366 material in total arsenic concentration in soils and make us believe that the main source of arsenic, in 367 the present study, is geogenic.

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Figure. 2. Raw data for total As concentrations in soils from different sites in the Amazon region for 373 the layers 0-20, 20-40 and 40-60 cm. Each site correspond to a Brazil nut stand, identified by: SM 374 (Sena Madureira, state of Acre); XP (Xapuri, state of Acre); ANO (Anori, state of Amazonas); AR 374 (Sena Madureira, state of Acre); XP (Xapuri, state of Acre); ANO (Anori, state of Amazonas); AR 375 (Aruanã farm-Itacoatiara, state of Amazonas); AP (Laranjal do Jari, state of Amapá); MT (Itaúba, state 376 of Mato grosso); PA (Santarém, state of Pará); RO (Porto Velho, state of Rondônia) and RR 377 (Caracaraí, state of Roraima). The letter on top of each state indicates the result of Tukey's HSD test (P $378 = 0.05$ on the log-transformed As concentration, following a significant ANOVA ($p < 0.001$). 379

380 **3.23 Extractable soil As (soluble and available)**

 Extractions for soluble and available As showed the highest concentrations for the site SM which is in accordance with the results observed for total As. For the soluble fraction the 383 concentrations ranged from <0.047 (MT) to 16 μ g kg⁻¹ (AC:SM). Available As varied from 2.14 (AM:AR) to 236.5 µg kg⁻¹ (AC:SM) (**Figure**. **3**). Observing the mean concentrations for soluble As, it is noteworthy that in SM the values are

386 increasing from the topsoil to the deepest layers. On the other hand, in the sites AM:AR, AM:ANO,

387 PA, RO and AP, soluble As decreases from the topsoil (0-20 cm) to the deepest layers (20-40 and 40-

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 60 cm). The sites RR and AP presented the lowest average concentration for soluble As (**Figure. 3**). The site with highest pH (AP) also presented the lowest median values for soluble As (**Figure. 3**). This effect is confirmed by the negative correlation between soluble As and pH in all the three layers studied (**Table 3**).

392 The average concentration for available As in AC:SM in the layers 0-20, 20-40 and 40-60 cm 393 (45.3, 68.6 and 117 μ g kg⁻¹) are at least 4-fold higher than those observed in all other sites studied. There was no meaningful variation among the sites for available As.

 The observed behavior of soluble and available As in these soils may be partly explained by their interactions with available P (**Table 2**), which showed positive correlation with both fractions in all the three layers studied (**Table 3**). This is expected as a greater content of P in oxidic soils (which retain phosphate strongly) make As much more available and soluble due to competitive adsorption on positively-charged sites (Campos et al., 2013a).

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403 **Figure 3.** Raw data for concentration of As available and soluble for 9 sites in the Amazon region.
404 | Each site correspond to a Brazil nut stand, identified by: AC-SM (Sena Madureira, state of Acre); Each site correspond to a Brazil nut stand, identified by: AC:SM (Sena Madureira, state of Acre); 405 | AC:XP (Xapuri, state of Acre); AM:ANO (Anori, state of Amazonas); AM:AR (Aruanã farm-
406 | Itacoatiara, state of Amazonas): AP (Laranial do Jari, state of Amapá): MT (Itaúba, state of Mato Itacoatiara, state of Amazonas); AP (Laranjal do Jari, state of Amapá); MT (Itaúba, state of Mato 407 grosso); PA (Santarém, state of Pará); RO (Porto Velho, state of Rondônia) and RR (Caracaraí, state 408 of Roraima). The letter on top of each state indicates the result of Tukey's HSD test ($p = 0.05$) on the 408 of Roraima). The letter on top of each state indicates the result of Tukey's HSD test ($p = 0.05$) on the 409 log-transformed As concentration, following a significant ANOVA ($p < 0.001$) using generalized least 409 log-transformed As concentration, following a significant ANOVA ($p < 0.001$) using generalized least 410 squares. squares.

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412 Considering the proportional contribution given by each fraction (soluble and available) to the 413 extractable As concentration, it is noticeable that available As represents the major fraction with more 414 than 70% of the As being extracted by KH_2PO_4 on average in all the sites studied. Generally, the 415 contribution given by the soluble fraction in all the sites reduces from the topsoil to the deepest layer, 416 as follows: 25% > 18% > 16% (0-20, 20-40 and 40-60 cm respectively) (**Figure. 4**). The site with the 417 greatest contribution given by available As was AC:SM with 92%, 94% and 94% for the 0-20, 20-40 418 and 40-60 soil layers. On the other hand, the sites with smallest contribution given by available As

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419 were RR in the layer 0-20 cm (67.7%), PA in the layer 20-40 cm (75.7%) and $\overrightarrow{AM}:AR$ in the layer 40-60 cm (77.5%) (**Fig.ure 4**).

 Arsenic availability was also assessed using tri-distilled water, ammonium sulfate, ammonium phosphate, ammonium oxalate + oxalic acid, a mixture of organic acids, Mehlich-1, and USEPA 3051A in artificially contaminated soils (entisol and oxisol) by Melo et al. (2012). The extracting solutions tested in the experiment were equally efficient in assessing plant availability of As in the soils, however there was greater concentration of As extracted from the clayey entisol when compared with the medium-textured oxisol due to differences between the adsorptive capacities of these soils. Besides validating extractants for plant-available As, these authors emphasized the capacity of tropical 428 soils, with variable charge, to adsorb As, especially those soils rich in clay (Melo et al., 2012).

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Figure 4. Relative contribution in % from soluble and available As in soils from different sites in the 433 | Amazon region. Each site correspond to a Brazil nut stand identified by: AC-SM (Sena Madureira. 433 Amazon region. Each site correspond to a Brazil nut stand identified by: AC:SM (Sena Madureira, 434 state of Acre); AC:XP (Xapuri, state of Acre); AM:ANO (Anori, state of Amazonas); AM:AR 434 State of Acre); AC:XP (Xapuri, state of Acre); AM:ANO (Anori, state of Amazonas); AM:AR
435 (Aruanã farm-Itacoatiara, state of Amazonas); AP (Laranial do Jari, state of Amapá); MT (Itaúba, state 435 (Aruanã farm-Itacoatiara, state of Amazonas); AP (Laranjal do Jari, state of Amapá); MT (Itaúba, state 436 of Mato grosso); PA (Santarém, state of Pará); RO (Porto Velho, state of Rondônia) and RR (Caracaraí, state of Roraima). 438

439 **3.43 Relationship between soil As and physicochemical attributes**

 Total As concentration (total As) correlated positively (p < 0.05) with CEC, available P, K, exchangeable Al, and silt in the three layers studied and also with available Zn in the 20-40 and 40-60 442 cm layers. Total As correlated negatively ($p < 0.05$) with pH, exchangeable Ca, available Cu in the three layers studied and also with exchangeable Mg and available Mn in the 0-20 cm layer and with available B and sand in the 0-20 and 20-40 cm layers. These results indicates that, in tropical soil conditions, increasing pH values will decrease total As content either directly because of its influence on colloidal charges (i.e., causing a decrease in positive charges) or indirectly influencing the

 availability and retention of competing anions such as phosphate and sulfate (considering that soils are open systems and available forms can be removed, thereby reducing total levels) (**Table 3**).

449 Soluble As correlated positively $(p < 0.05)$ with CEC, available P, exchangeable Al and silt in the three layers evaluated and also with OM and clay in the 0-20 cm layer and with available Zn and K 451 in the 40-60 cm layer. Soluble As correlated negatively $(p < 0.05)$ with pH, exchangeable Ca and available Cu in the three layers studied and also with exchangeable Mg, available B and sand in the 0- 20 and 20-40 cm layers following by Fe in the 40-60 cm and available Mn in the 0-20 and 40-60 cm layers (**Table 3**).

 On the other hand, available As correlated positively (p < 0.05) with available P, K, Zn, exchangeable Al and silt in all the three layers studied and also with CEC in the 20-40 and 40-60 cm layers and with available Mn in the 20-40 cm layer following by sand in the 40-60 cm layer. Available As correlated negatively (p < 0.05) with OM and clay in the 20-40 and 40-60 cm, and also with pH in the 20-40 cm, with extractable S in the 0-20 and 40-60 cm layers, Fe in the 40-60 cm layer and available Cu in the 0-20 cm layer (**Table 3**).

 In the present study, an important factor influencing As bioavailability that negatively affected total and soluble As was soil pH as can be observed in the figure 5. As pointed out by Zhang et al. (2017), numerous studies have shown that soil pH is an important factor controlling arsenic 464 geochemistry, with adsorption of As decreasing as pH increases. The decrease in As bioavailability when the pH is enhancedis is attributed to the higher amounts of negative charges on soil components that generate variable charge, such as 1:1 clays, along with less amounts of positive charges on Fe and Al oxides. Such pH influence on charge generation is pronounced in tropical soils, since these are 468 variable charge soils (Fontes et al., 2001).

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474 According to Huang et al. (2006), NaH₂PO₄-extractable As correlates well with total As concentrations found in edible parts of various crops and soils collected from paddy rice fields in suburban areas of Fukan province in southeast China, showing the usefulness of this extractant to assess the bioavailability of As in soils. These authors also reported that As availability decreased significantly with increasing silt size (0.02–0.002 mm) and free iron content, but increased significantly with soil pH and organic matter content. Those findings are not supported by results observed in the present study for the silt fraction, which showed positive correlation with total, soluble, and available As in soil samples of all the three layers studied (**Table 3**).

 Similarly to what was observed in the present study (20-40 and 40-60 cm layers), Melo et al. (2012) observed that As availability (expressed by concentration of As in plants) was negatively correlated with clay content. The main reason for finding negative correlations between available As and clay content, especially in subsurface, is because of the particular composition and characteristics of the clay fraction in tropical soils, which presents significant amounts of Fe and Al oxides and oxyhydroxides as well as high ZPC (Fontes et al., 2001). According to McBride (1994), chemical adsorption of As occurs mainly in Fe and Al oxides, as well as in amorphous aluminosilicates, with small quantities being retained by clay silicates and organic substances. For this reason, the mobility and bioavailability of As is generally lower in clayey soils with variable charge.

 It is noteworthy that the highest correlation coefficients were observed between exchangeable Al $(A³⁺)$ and total, available, and soluble As in all the three layers studied, which were all positive (*p* $493 \leq 0.05$) (**Table 3**). The mechanism by which high Al³⁺ is associated with high As concentrations may be related to the preferential retention of Al in the negative sites of variable-charge minerals (including clay minerals, Fe and Al oxides and oxyhydroxides), thus creating a bridge for binding As through formation of surface ternary complexes, which has been previously reported (Oliveira et al., 2015).

 Given that As (as arsenate) and P (as phosphate) present similar chemical behavior in soils, available P has an antagonistic effect upon As retention, as P can compete with As for anionic sorption sites (Sharma et al., 2011). This explains the positive correlation observed between the soluble and available As fractions and the available P in soil samples of all the three layers studied (**Table 3**). Several studies show evidence supporting a strong relationship between As availability in soil and soil available P as a consequence of the chemical and structural similarities between their anionic forms (Punshon et al., 2017; Arco-Lázaro et al., 2018; Chowdhury et al., 2018). Jiang et al. (2014) observed significant positive correlation between soil available P and exchangeable As. The authors also stated that the main factors influencing As uptake by rice in that study included soil pH, total organic carbon,

 as well as soil available P and Fe. In this case, soil pH affected As uptake indirectly via impact on available P and Fe. Wang et al. (2018) conducted a pot experiment to study the effect of P application on the lability and phytoavailability of As in the soil and observed that application of exogenous phosphate induced soil arsenic mobilization, thereby increasing arsenic lability due to a competition effect. Finally, a positive correlation found between available P and total As in all soil layers is 511 eertainly probably related to the fact that positively-charged colloidal surfaces that adsorb P in tropical agroecosystems are also relevant for retaining As in the soil system.

 Extractable sulfur in this study affected negatively the amount of available As in the soil, in the layers 0-20 and 40-60 cm. The competition between S and As, mainly in the 40-60 cm layer, where sulfur could have more affinity with the binding sites in tropical soils is relevant if we consider \parallel that As in soil solution cannot overcome S informed the sites available, due to a concentration effect, i.e., S in solution is generally higher than As in solution in most soils. Therefore the decreasing of available As in a long term is certain when the concentration of S is higher. But this is not the behavior often showed in experiments of shorter term where the addition of S displace As from the binding sites and increase its concentration in soil solution. In fact, Song et al. (2015) studying the effect of sulfate on As mobility in contaminated soils observed that addition of sulfate significantly enhanced As availability. Moreover, in a study aiming to evaluate the competitive adsorption of arsenate/phosphate and arsenate/sulfate in Brazilian soils, Campos et al. (2013a) observed that the addition of phosphate and sulfate decreased arsenate adsorption and consequently increased the concentration of As in soil solution and its availability.

 The negative correlation between OM and available As in the 20-40 and 40-60 cm layers observed in the present study (**Table 3**), suggest that OM can be considered trap and should be able to alleviate As mobility in soil by formation of covalent bindings, especially in OM-rich soils which is 529 the case of some points sampled in these sites studied (Langner et al., 2012).

 The adsorption capacity of As by soils provides an indication of the system's buffer power, since many soil attributes (such as texture, Fe and Al oxide content, among others) interfere with the buffer capacity and, consequently, influence leaching losses (Campos et al., 2013a). Although we did not find any relevant correlation among Fe and As fractions in soils from the present study, a great number of studies in the literature shows otherwise. In fact, a close coupling between the biogeochemical cycles of iron and arsenic in both oxidizing and reducing environment has been well established (Dixit and Hering, 2003). The mechanism that is presumably involved in this interaction is that both As(III) and As(V) adsorb to iron (hydr) oxides through, in part, formation of inner-sphere complexes (Tufano et al., 2008). Wang et al. (2012), studying As retention in muddy sediments observed that As retention was strongly coupled to Fe and S, and correlation analysis showed that total and extractable As correlated positively and well with total Fe and HCl-extractable Fe.

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542 **Table 3**. Pearson's correlation coefficients (r) for log transformed physicochemical variables versus 543 total, soluble and available arsenic concentrations from different sites in the Amazon region in the 543 total, soluble and available arsenic concentrations from different sites in the Amazon region in the layers 0-20 cm, 20-40 and 40-60 cm.

Soil	$---0-20$ cm $---$			$---20-40$ cm $---$			-40-60 cm----- $- - - -$		
attribute	As	$As-$ KNO ₃	$As-$ KH_2PO_4	As	$As-$ KNO ₃	$As-$ KH_2PO_4	As	$As-$ KNO ₃	$As-$ KH_2PO_4
pH	$-0.43**$	$-0.33**$	-0.14	$-0.44**$	$-0.48**$	$-0.17*$	$-0.46**$	$-0.36**$	-0.17
OM	0.06	$0.27**$	-0.12	-0.02	0.11	$-0.19*$	-0.12	-0.15	$-0.23**$
CEC	$0.34**$	$0.26**$	0.15	$0.33**$	$0.35**$	$0.22*$	$0.39**$	$0.39**$	$0.30**$
P	$0.27**$	$0.39**$	$0.43**$	$0.31**$	$0.25**$	$0.32**$	$0.29**$	$0.23**$	$0.31**$
S	0.03	0.08	$-0.23*$	0.14	0.10	-0.14	0.06	-0.05	$-0.19*$
Al	$0.53**$	$0.40**$	$0.41**$	$0.52**$	$0.55**$	$0.44**$	$0.56**$	$0.60**$	$0.47**$
Fe	0.08	0.09	-0.01	-0.06	0.09	-0.15	-0.15	$-0.19*$	$-0.29**$
Clay	0.15	$0.26**$	-0.06	0.04	0.10	$-0.21*$	0.01	0.0	$-0.29**$
Silt	$0.55**$	$0.31**$	$0.40**$	$0.57**$	$0.45**$	$0.37**$	$0.56**$	$0.51**$	$0.40**$

545 Soil attributes: total As (HF digestion); soluble arsenic (KNO3-extractable As); available arsenic 546 (KH₂PO₄-extractable As); pH (pH in H₂O); P (available phosphate); S (extractable S-SO₄²); CEC 547 (cation exchange capacity); OM (organic matter); Al (exchangeable aluminum); Fe (available iron); 548 clay and silt $(\frac{9}{6})$. * Significant correlation (p < 0.05); ** significant correlation (p < 0.01). The 549 correlation includes all the sites studied except the site "Santarém/PA" because the data for texture was obtained in only one point. was obtained in only one point. 551

552 **Figure 6** shows the gradient of influence of the soil chemical variables and correlation groups 553 of the samples for each site studied in the Amazon region. The model including the two principal 554 components accounted for 52.7% of the total variance explained in the three layers studied (29.8% for

555 PC1 and 22.9% for PC2).

 The first component (PC1) was negatively dominated by P, S, K, total As, available As, soluble As, CEC, OM, Al and Fe and positively dominated by Zn, Mn, Cu, B, Mg, Ca and pH, which accounted for 29.8% of the variance. The second (PC2) was negatively dominated by S and positively dominated by Fe, OM, Al, AS, CEC, soluble As, available As, K and P, representing 22.9% of the

560 total variance in the data (**Figure. 6**).

 The soil attribute with the biggest contribution for PC1 is soil pH, followed by Ca and Mg, i.e., in terms of acidity and the presence of basic cations, soils from the Amazon are very diverse, reflecting a heterogeneous mineralogy and capacity to adsorb anions as can be observed by the groupings of sites formed along the axis of these attributes. The soil attribute that contributed the most for PC2 is Zn, followed by K and P, which are influencing the grouping of the samples from the site XP.

 Total, soluble, and available As are positively correlated with K, P, CEC, OM, Al and Fe and 568 also present a major influence in the samples of the sites AM:ANO, AC:XP and AC:SM. The close correlation between the As fractions and CEC and OM suggest that these soil properties are important factors to retain these As forms in the soil. With regard to CEC, unless there is a formation of a ternary complex in the interface soil-solution with an intermediate cation in the system, we believe that the correlation between CEC and As fractions is indirectly influenced by soil organic matter. The extractable S was separated from the other soil chemical attributes in the third quadrant and showed no relationship with total and soluble As, and only a weak relationship with available As (negative correlation in the 0-20 and 40-60 cm layers) (**Table 3**) and therefore represented a major influence for the samples from the sites AM:AR and PA (**Figure. 6**).

 The site AP forms a separated group in the border of first and fourth quadrant, which indicates that this site present its chemistry distinct from the others, mainly because of the higher values of pH, Ca, Mg, B, Mn, and Cu. On the other hand, this site receives lower impact from the total, soluble and available As fractions because of its lower As concentrations compared with the other sites (**Figure. 6**). In summary, soil pH in combination with CEC and OM might be the key factors that influence As concentrations in soil as well as the concentrations of soluble and available As in distinct soil layers (depths) for the conditions of the present study.

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592 farm-Itacoatiara, state of Amazonas); AP (Laranjal do Jari, state of Amapá); MT (Itaúba, state of Mato 593 grosso); PA (Santarém, state of Pará); RO (Porto Velho, state of Rondônia) and RR (Caracaraí, state 593 grosso); PA (Santarém, state of Pará); RO (Porto Velho, state of Rondônia) and RR (Caracaraí, state 594 | of Roraima). The symbols represent each soil depth as follow – circle: 0-20 cm; cross: 20-40 cm and 594 of Roraima). The symbols represent each soil depth as follow – circle: 0-20 cm; cross: 20-40 cm and 595 triangle: 40-60 cm. Soil chemical attributes: S (extractable S-SO $_4^2$ in mg kg⁻¹); available Mn, Zn, Fe, 595 $\frac{t}{\text{triangle: }40\text{-}60 \text{ cm.}}$ Soil chemical attributes: S (extractable S-SO₄² in mg kg⁻¹); available Mn, Zn, Fe, 596 Cu, B, K, and P (mg kg^{-1}); exchangeable Mg, Ca, and Al in cmol_c dm⁻³; CEC (cation exchange 597 | capacity in cmol_c dm⁻³); OM (organic matter in %); pH (pH in H₂O); total As (mg kg⁻¹); soluble As 598 (μ g kg⁻¹); available As (μ g kg⁻¹). 599

600 **4. CONCLUSIONS**

601 Soil pH and CEC indirectly are the main chemical attributes influencing the dynamic of As

602 fractions in soils from the sites studied in the Amazon basin. In a smaller magnitude but not less

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